

1 **The Fractional MacroEvolution Model:**

2 **A simple quantitative scaling macroevolution model**

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20 **Abstract:**

21 Scaling fluctuation analyses of the marine animal diversity, extinction and origination
22 rates based on the Paleobiology Database occurrence data have opened new perspectives on
23 macroevolution, supporting the hypothesis that the environment (climate proxies) and life
24 (extinction and origination rates) are scaling over the “megaclimate” biogeological regime
25 (from ≈ 1 Myr to at least 400 Myrs). In the emerging picture, biodiversity is a scaling “cross-
26 over” phenomenon being dominated by the environment at short time scales and by life at long
27 times scales with a cross-over at ≈ 40 Myrs. These findings provide the empirical basis for
28 constructing the Fractional MacroEvolution Model (FMEM), a simple stochastic model
29 combining destabilizing and stabilizing tendencies in macroevolutionary dynamics, driven by
30 two scaling processes: temperature and turnover rates.

31 Macroevolution models are typically deterministic (albeit sometimes perturbed by
32 random noises), and based on integer ordered differential equations. In contrast, the FMEM is
33 stochastic and based on fractional ordered equations. Stochastic models are natural for systems
34 with large numbers of degrees of freedom and fractional equations naturally give rise to scaling
35 processes.

36 The basic FMEM drivers are fractional Brownian motions (temperature, T) and fractional
37 Gaussian noises (turnover rates E_+) and the responses (solutions), are fractionally integrated
38 fractional Relaxation processes (diversity (D), extinction (E), origination (O) and $E_- = O - E$).
39 We discuss the impulse response (itself a model for impulse perturbation, e. g. bolide impacts)
40 and derive the full statistical properties including cross covariances. By numerically solving
41 the model, we verified the mathematical analysis and compared both uniformly and irregularly
42 sampled model outputs to paleobiology series.

43 **1. Introduction**

44

45 Several centuries of paleontological research revealed that the evolution of Life on Earth
46 is characterized by high temporal complexity characterized by periods of sluggish and
47 predictable evolution (Jablonski, 1986; Casey et al., 2021) with mass extinctions characterized
48 by selectivity that is low or different in kind than in “background intervals” (Raup, 1992a ;
49 Raup, 1994 ; Payne & Finnegan, 2007). There are also mass evolutionary radiations which
50 sometimes are contemporaneous with mass extinctions (Cuthill et al., 2020). Moreover
51 apparently the factors and modes of macroevolution vary with time—e.g. Cambrian explosion
52 or Ediacaran-Cambrian radiation and post-Cambrian evolution (Gould, 1990 ; Erwin, 2011 ;
53 Mitchell et al., 2019); environment (Kiessling et al., 2010; Jablonski et al., 2006; Miller &
54 Foote, 2009 ; Boyle et al., 2013; Spiridonov et al., 2015; Tomašových et al., 2015); and
55 timescales (Crampton et al., 2018 ; Van Dam et al., 2006 ; Spiridonov et al., 2017b; Beaufort
56 et al., 2022). Moreover macroevolution is strongly influenced by Earth system —geological,
57 climatic, and paleoceanographic—factors (Marshall et al., 1982; Lieberman & Eldredge, 1996;
58 Lieberman, 2003 ; Saupe et al., 2019; Halliday et al., 2020; Carrillo et al., 2020), but also by
59 biotic interactions, which can translate into patterns which are apparent on extremely long time
60 scales of tens to hundreds of millions of years (Vermeij, 1977 ; Jablonski, 2008 ; Erwin, 2012;
61 Vermeij, 2019). Also, there are questions on the role of general stochasticity and path
62 dependence/memory in evolutionary dynamics (Schopf, 1979; Hoffman, 1987; Erwin, 2011 ;
63 Erwin, 2016 ; Gould, 2001; Gould, 2002 ; Cornette & Lieberman, 2004). The question is:
64 can we reconcile in a single simple model this multitude of hierarchically organized and
65 causally heterogenous processes producing macroevolutionary dynamics, while maintaining
66 simplicity and conceptual clarity? Here we argue that we can.

67 The development of large, high temporal resolution databases – both of past climate
68 indicators (Veizer et al., 1999 ; Song et al., 2019 ; Grossman & Joachimski, 2022) and of
69 paleobiological information such as Paleobiology Database (Alroy et al., 2001 ; Alroy et al.,

70 2008) or NOW (Jernvall & Fortelius, 2002; Žliobaitė et al., 2017; Žliobaitė, 2022), is
 71 transforming our understanding of macroevolution. Time series are frequently long enough
 72 that they be studied systematically - not just as chronologies to be compared with other
 73 chronologies - but as functions of temporal *scale*, i.e. the behaviour of their fluctuations as
 74 functions of duration (or equivalently, their behaviour as functions of frequency). A regime
 75 over which fluctuations ΔT are scaling i.e. of the form $\Delta T(\Delta t) \propto \Delta t^H$ where Δt is duration - “lag”,
 76 scale, and H is an exponent - can be used to objectively define dynamical regimes (this scaling
 77 relationship holds in a statistical sense discussed below). This is because over such a regime,
 78 long duration fluctuations at scale $\lambda \Delta t$ ($\lambda > 1$) are related to the shorter duration fluctuations
 79 by: $\Delta T(\lambda \Delta t) = \lambda^H \Delta T(\Delta t)$ i.e. the fluctuations at different time scales differ only in their
 80 amplitudes. In addition, we can already distinguish the qualitatively different types of regime
 81 by the sign of the exponent H . $H > 0$ implies that fluctuations increase with scale whereas $H < 0$
 82 implies that they decrease.

83 An important consequence for our understanding of deep time biogeodynamics - here
 84 understood as joint Earth-Life systems - is the robustness of the “megaclimate” regime of
 85 positively scaling (a short hand for $H > 0$) with time scale temperature fluctuations meaning that
 86 at longer time scales climates become more and more distinct, first (Lovejoy 2013), (Lovejoy
 87 2015) on the basis of long paleotemperature data from ocean core stacks (Veizer et al. 2000),
 88 (Zachos et al. 2001). Megaclimate is the hypothesis that there is a unique (presumably highly
 89 nonlinear) biogeological dynamical regime that operates over time scales spanning the range
 90 ≈ 1 Myr to (at least) several hundred Myrs. This would be the consequence of a unique (albeit
 91 complex, nonlinear) underlying dynamic that is valid over this wide range of scales;
 92 presumably it involves a scaling (hence hierarchical) mechanism that operates from long to
 93 short durations. A consequence is the existence of a statistical scaling regimes (notably of
 94 paleo temperatures), empirically verified throughout the Phanerozoic. While its inner scale

95 appears to be fairly robust at around 1 Myr, its outer scale (the longest duration over which it
96 is valid) is not known although it appears to be at least 300 Myrs. The megacclimate regime
97 implies that the underlying biology - climate dynamics are essentially the same over these time
98 scales: i.e. that the statistics are stationary (although they may *appear* to be nonstationary at
99 shorter time scales).

100 The hypothesis that biology and the climate are linked, and that climate is crucial and
101 defining variable in ecological and evolutionary turnovers (Vrba, 1985 ; Vrba, 1993 ; Eldredge,
102 2003; Lieberman et al., 2007; Hannisdal & Peters, 2011; Mayhew et al., 2012 ; Crampton et
103 al., 2016; Spiridonov et al., 2016; Spiridonov et al., 2017a; Spiridonov et al., 2020a;
104 Spiridonov et al., 2020b; Mathes et al., 2021), is hardly controversial - after all - the
105 paleoclimate indicators themselves are often based on stable oxygen isotopic analyses of
106 CaCO₃ from ancient foraminifera (O'Brien et al., 2017) or sometimes for more recent periods
107 estimated directly from occurrences and abundances of taxa using modern analog techniques
108 (Dowsett & Robinson, 1998; Green, 2006). However, the scope and utility of the megacclimate
109 notion would increase if it could be backed up by direct analysis of paleobiological series,
110 particularly of extinction and origination rates. This has now been done. A recent paper
111 (Spiridonov and Lovejoy 2022), hereafter SL) found that genus - level extinction and
112 origination rates exhibited scaling statistics over roughly the same range as the paleo
113 temperatures confirming that the megacclimate includes these key macroevolutionary
114 parameters.

115 The shortest scale of SL's paleobiological time series was closer to ≈ 3 Myrs (average
116 stage resolution was 5.9 Myrs) which correspond to the durations of shortest Paleobiology
117 Database stages – a standard shortest time resolution for Phanerozoic scale global biodiversity
118 analyses (e.g. (Alroy et al., 2008; Alroy, 2010b)). Systematic reviews and multiple case studies
119 revealed that even variously defined (molecular, morphological, phylogenetic, and taxic)

120 evolutionary rates universally exhibit negative time scaling behavior (Gingerich, 1993;
121 Gingerich, 2001; Gingerich, 2009; Roopnarine, 2003; Harmon et al., 2021; Spiridonov &
122 Lovejoy, 2022), which suggests the universality of the temporal scaling - hence hierarchical -
123 evolutionary dynamics. Although an inner megacclimate scale of ≈ 1 Myrs was also proposed
124 in (Lovejoy 2013), (Lovejoy 2015) and is discussed at length in the nonspecialist book
125 (Lovejoy 2019). The scaling, and thus by implication dominance of time symmetric
126 hierarchical interactions, was also detected on multimillion year time scales in sedimentation
127 rates/stratigraphic architecture (Sadler, 1981), sea level (Spiridonov and Lovejoy, 2022), and
128 dynamics of continental fragmentation (Spiridonov et al., 2022), which shows universality of
129 the pattern in major Earth systems as well. Therefore, the time scaling patterns of evolution
130 and megacclimate overlap at the very wide range of temporal scales (from $\approx 10^6$ to $> 4 \times 10^8$ yrs),
131 which motivates the development of quantitative models which explicitly tackle and integrate
132 together these time scale symmetries.

133 If macroevolution and climate respect wide range scaling, then it may be possible to
134 resolve a longstanding debate in macroevolution. In terms first posed by (Van Valen 1973),
135 we may ask: are evolutionary processes dominated by external factors - especially climate, the
136 “Court Jester” (Barnosky, 2001; Benton, 2009) - or is life itself - the “Red Queen” (Van Valen,
137 1973 ; Finnegan et al., 2008) - the determining process. SL proposed a scaling resolution of
138 the debate in which at scales below a critical transition time τ of ≈ 40 Myrs, the climate process
139 is dominant, but there is a “cross-over” beyond which life (self-regulating by means of
140 geodispersal and competition) are dominant. SL thus quantitatively concluded that at long
141 enough time scales the Red Queen ultimately overcomes the Court Jester. The scaling
142 processes of the Earth system here are playing double role (thus Geo-Red Queen theory) –
143 climate fluctuations growing with time scale cause perturbations in diversity to grow in their
144 size, but at the same time, at longer and longer time scales fluctuating climates and plate

145 tectonics cause the mixing and competitive matching of biota, thus effectively globally
 146 synchronizing it. Later results in a described “cross-over” when unstable and wandering
 147 diversity regime changes to longer time scale fluctuation canceling or stabilizing regime
 148 (*Spiridonov and Lovejoy, 2022*).

149 Physicists use the term “cross-over”, as a short-hand to describe analogous phenomena
 150 involving processes that are subdominant over one scale range but eventually become dominant
 151 at longer scales. However, such transitions are typically modelled by Markov processes so that
 152 the autocorrelations are exponential so that at the critical time scale, the transition between two
 153 regimes is fairly sharp. In SL, on the contrary, in keeping with the basic megacclimate scaling
 154 dynamics, the cross-over was postulated to be a the consequence of the interaction of two
 155 scaling processes i.e. the transition is a slow, power law one. An analogous scaling cross-over
 156 phenomenon was found in phytoplankton where the competing scaling processes were
 157 phytoplankton growth (with turbulence) and a predator-prey process of zooplankton grazing
 158 (*Lovejoy et al. 2001*).

159 SL argued that while both macro evolution and climate respect wide range statistical
 160 scaling, that their quantitative and qualitative differences are significant and this was the key
 161 to macroevolution power law cross-overs. While temperature (T) fluctuations vary with time
 162 scale Δt as $\Delta T(\Delta t) \approx \Delta t^{H_T}$ with $H_T \approx 0.25$, the corresponding laws for extinction (E) and
 163 origination (O) have $H_E, H_O \approx -0.25$. When $H > 0$, fluctuations grow with scale so that the
 164 corresponding series tend to “wander” without any tendency to return to a well-defined value,
 165 they appear “unstable”. On the contrary, when $H < 0$, successive fluctuations tend to have
 166 opposite signs so that they increasingly cancel over longer and longer time scales, they
 167 fluctuate around a long term value, they appear stable.

168 To deepen our understanding, it is necessary to build a quantitative model of the
 169 interaction of climate and life. In recognition of the strongly nonlinear nature of evolutionary

170 dynamics, there have developed numerous deterministic chaos models such as predator - prey
171 models (e.g. (Huisman and Weissing 1999), (Caraballoa et al. 2016)). Although extensions
172 with some stochastic forcing exist (e.g. (Vakulenko et al. 2018)), in the latter, the stochasticity
173 is a perturbing noise on an otherwise deterministic system. In paleontology the model of
174 exponential (unconstrained) proportional growth of diversity was historically popular (Stanley,
175 1979; Benton, 1995), or expanded for possible acceleration due to niche construction effects
176 (second-order positive feedback) - a hyperbolic model (Markov and Korotayev, 2007). These
177 simple models of expansion were contrasted by single or coupled logistic models of resource
178 constrained competitive macroevolutionary dynamics, sometimes also supplemented with
179 random perturbations which account for effects of mass extinctions (Sepkoski 1984; 1996); or
180 implicitly hierarchical, and also competition constrained Gompertz models (Brayard et al.,
181 2009). However, such models assume that only a few degrees of freedom are important
182 (typically fewer than 10) whereas the true number is likely to be astronomical. It is therefore
183 logical to model the process in a stochastic framework (involving infinite dimensional
184 probability spaces), where the primary dynamics are stochastic using the scaling symmetry as
185 a dynamical constraint. Therefore, there is growing recognition of stochastic models as
186 essential tools for understanding macroevolutionary dynamics. Actually some of the first
187 models that tried to explain complexities of macroevolutionary dynamics were stochastic
188 Markovian birth and death models (Raup, 1985 ; Raup & Valentine, 1983 ; Gould et al., 1977 ;
189 Raup, 1992a ; Nee, 2006). Several recent applications of linear stochastic differential equations
190 were used in causal inference of macroevolutionary drivers and competitive interactions
191 between clades (Reitan & Liow, 2017 ; Liow et al., 2015 ; Lidgard et al., 2021).

192 Beyond the realism of implicitly involving larger numbers of degrees of freedom,
193 stochastic models have the advantage that they may be linear even though the corresponding
194 deterministic model may be highly nonlinear. Also, by the simple expedient of using fractional

195 ordered differential equations rather than the classical integer ordered ones, stochastic models
196 can readily handle scaling which is rarely explicitly considered in macroevolutionary analyses.
197 This is because fractional equations have impulse response functions (Green's functions) and
198 hence solutions that are based on scaling (power laws) rather than the exponential Green's
199 functions associated with integer ordered differential equations.

200 In this paper, we therefore build a simple model for biodiversity (D) that can reproduce
201 and explain SL's findings. The model is parsimonious: has only two scaling drivers - the
202 climate and life – and by construction - it reproduces the observed scaling cross-over at
203 40Myrs. Although the model has two basic exponents (climate and life) and two correlation
204 coefficients, it satisfactorily reproduces the fluctuation statistics of D , T , E , O as well as the
205 turnover ($E_+ = O + E$) and difference $E_- = O - E$ over the range ≈ 3 Myrs to several hundred
206 Myrs (the longest scales available). Beyond this, it explains the 15 pairs of (scale by scale
207 fluctuation correlations) over the same observed range. The data are from SL paper-they
208 represent stage level time series of Phanerozoic marine animal genera O and E (second-for-
209 third origination and extinction proportion (Kocsis et al., 2019 ; Alroy, 2015) not-standardized
210 for the duration of stages), sample standardized using shareholder quorum method (Alroy,
211 2010a) D of Phanerozoic marine animals based on Paleobiology Database data
212 (<https://paleobiodb.org/>). While paleotemperatures (T) are also the same as in the SL paper,
213 obtained from (Song et al., 2019)

214 As a final comment, we should note that the basic – simplest - stochastic “cross-over”
215 process is the fractionally integrated fractional relaxation noise (ffRn process) whose properties
216 were only fully elucidated very recently (Lovejoy 2022) in the context of long term weather
217 forecasts (Del Rio Amador and Lovejoy 2021) and climate projections (Procyk et al. 2022).
218 The new model has conceptual commonalities with the environmental “stress model” of M.
219 Newman that attempted to replicate the scaling statistics of extinction intensities of marine

220 biota (Newman, 1997; Newman & Palmer, 2003). The model presented here is more
 221 sophisticated since it ties the principal macroevolutionary variables — O, and E — to a
 222 principal geophysical scaling process — the megacclimate — in producing realistic multi time
 223 scale global dynamics of marine animal biodiversity, while keeping its conceptual simplicity
 224 in transparently using a few crucial parameters of time scaling and correlations. The model also
 225 explicitly hierarchical through scaling relations – having a desirable feature of a unified
 226 evolutionary theory (Eldredge, 1985 ; Eldredge, 1989 ; Gould, 2002 ; Lieberman et al., 2007).
 227

228 **2. The model:**

229 2.1. The equations:

230 *2.1.1 The basic diversity equation*

231 The SL picture is one where the extra-biological factors (“the climate”) are scaling and
 232 drive biodiversity from $\approx 1\text{Myr}$ to $\approx 40\text{Myrs}$, where the cross-over occurs followed by the
 233 domination of biotic-regulation at the longer time scales, which also enabled by global
 234 homogenization of biota at long time scales by continental drift and changes in climate zones
 235 (Geo-Red Queen dynamics). Based on this picture, we propose the following Fractional Macro
 236 Evolution Model (FMEM). At first we describe the model, we then comment on it.

237 The basic diversity equation is:

$$238 \quad \tau^h \frac{d^h(D - s_T T)}{dt^h} + D = s_E E_+; \quad E_+ = O + E$$

239 (1)

240 τ is the cross-over time scale ($\approx 40\text{Myrs}$) and $E_+ = E + O$ is the turnover rate. Whereas
 241 D , E_+ are already nondimensional, T must be nondimensionalized, for example by the standard
 242 deviation of its increments at some convenient reference scale, say 1 Myr. s_T , s_E are constants
 243 that are determined by the coupling between T and D (s_T) and E_+ and D (s_E).

244

245 *2.1.2 The drivers:*

246 The basic drivers are the climate (T) and life (E_+), themselves driven by Gaussian white
 247 noises γ_T, γ_E :

$$\begin{aligned} \tau^{\alpha+h} \frac{d^{\alpha+h} T}{dt^{\alpha+h}} &= \gamma_T \\ \tau^\alpha \frac{d^\alpha E_+}{dt^\alpha} &= \gamma_E \end{aligned} \quad (2)$$

248

249 α is the basic biology (extinction and origination rate) exponent ($\alpha \approx 0.25$ as deduced
 250 from SL's analysis) and h is the exponent difference (contrast) between the temperature and
 251 biology, from SL's analysis $h = 0.75 - \alpha \approx 0.5$. Combined with the diversity equation (eq. 1),
 252 these determine D . The derivatives are fractional, in this paper we use the semi-infinite "Weyl"
 253 fractional derivatives. For the arbitrary function $W(t)$, the ζ ordered Weyl fractional derivative
 254 is defined as:

$$\frac{d^\zeta W}{dt^\zeta} = \frac{1}{\Gamma(1-\zeta)} \frac{d}{dt} \int_{-\infty}^t (t-s)^{-\zeta} W(s) ds; \quad 0 < \zeta < 1 \quad (3)$$

255

256 Since fractional derivatives (and their inverse, fractional integrals) are – as in eq. 3 –
 257 generally convolutions, different fractional operators are defined on different ranges of
 258 integration for the convolutions. Weyl derivative are particularly simple to deal with since they
 259 are simply power law filters in Fourier space, see below (see e.g. (Miller and Ross 1993),
 260 (Podlubny 1999) for more information on fractional equations).

261

262 The γ 's are Gaussian white noises, they are proportional to "unit" white noises γ . Unit
 263 white noises have the properties:

$$\langle \gamma(t_1) \gamma(t_2) \rangle = \delta(t_1 - t_2); \quad \langle \gamma^2 \rangle = 1; \quad \langle \gamma \rangle = 0 \quad (4)$$

263

264 where the angle brackets indicate ensemble (statistical) averaging. Eq. 2 therefore implies that
 265 T , E_+ are fractional integrals of white noises. Depending on the value of the exponents), these
 266 are fractional Gaussian noises (fGns) and fractional Brownian motions (fBms), (Mandelbrot
 267 and Van Ness 1968) (see the later discussion on the small and large scale limits).

268

269 *2.1.3 Closing the model, the Diagnostic Equation:*

270 The preceding equations 1, 2 determine D , E_+ , T . However, in order for the model to
 271 determine E and O , we need a final equation for E_- :

$$272 \quad E_- = \tau_D \frac{dD}{dt} ; \quad E_- = O - E \quad (5)$$

273 This is just the differential form of the usual discrete - time definition of diversity:

274 $D_{j+1} = D_j (1 + O_j - E_j)$ where j is a time index. τ_D is the discretization time, it is the basic
 275 resolution of the series. Equation 5 plays no role in the dynamics, conventionally, it is the
 276 definition of D . Mathematically, eq. 5 is thus a "diagnostic equation" because it simply allows
 277 us to close (complete) the model by determining O , E :

$$278 \quad \begin{aligned} O &= (E_+ + E_-) / 2 \\ E &= (E_+ - E_-) / 2 \end{aligned} \quad (6)$$

279 *2.2 Discussion:*

280 *2.2.1 Diversity as a Fractionally Integrated Fractional Relaxation (ffRn) process*

281 The diversity model was written in a nonstandard way (eq. 1) because in this form, it's
 282 basic behaviour is transparent. When $h > 0$, the fractional term is the highest order derivative,
 283 at high frequencies it therefore dominates the zeroth order (D) term so that at short lags $\Delta t < \tau$,
 284 diversity fluctuations $\Delta D \propto \Delta T$ so that D follows the temperature. However at low frequencies
 285 ($\Delta t > \tau$), the zeroth order term dominates and we have instead $\Delta D \propto \Delta E_+$. By inspection, the

286 model therefore reproduces the cross-over at lag τ , and the crossover will be scaling due to the
 287 scaling of T , E_+ (eq. 2). The mathematical structure of the model is clearer if we substitute the
 288 driver in terms of their own Gaussian forcings γ_T , γ_E (eq. 2), rewriting eq. 1 as:

$$289 \quad \tau^h \frac{d^h D}{dt^h} + D = \tau^{-\alpha} \frac{d^{-\alpha}}{dt^{-\alpha}} (s_T \gamma_T + s_E \gamma_E) \quad (7)$$

290 $(d^{-\alpha} / dt^{-\alpha}$ is a fractional integral order α : for Weyl derivative and integrals it is the
 291 inverse of the α order derivative d^α / dt^α).

292 The linear combination of white noises $s_T \gamma_T + s_E \gamma_E$ is also a white noise. The D
 293 equation, is thus an order h fractional relaxation equation forced by an order α fractionally
 294 integrated white noise, i.e. it is a “fractionally integrated fractional relaxation” process (ffRn,
 295 (Lovejoy 2022)). The basic “unit” ffRn process $U_{h,\alpha}(t)$ satisfies:

$$296 \quad \left(\frac{d^{h+\alpha}}{dt^{h+\alpha}} + \frac{d^\alpha}{dt^\alpha} \right) U_{\alpha,h} = \gamma \quad (8)$$

297 Where γ is the unit white noise defined above and we have used the fact that for Weyl
 298 fractional derivatives fractional differentiation and integration commute. If time is rescaled
 299 ($t \rightarrow t/\tau$), we see (from eq. 7) that D is proportional to $U_{\alpha,h}$. We note that if $h = 1$, the D equation
 300 (eq. 1) would be a classical relaxation equation and if forced by a white noise (i.e. if $\alpha = 0$), D
 301 would be a classical Ornstein-Uhlenbeck (OU) process. OU processes are currently
 302 conventional approaches to the modeling and analysis of microevolutionary as well as
 303 macroevolutionary dynamics (Khabbazian et al., 2016; Bartoszek et al., 2017 ; Liow et al.,
 304 2022).

305

306 *2.2.2 Deterministic behaviour: Impulse response functions*

307 The D process - the solution to eq. 7 - is the response of the operator $\left(\frac{d^{h+\alpha}}{dt^{h+\alpha}} + \frac{d^\alpha}{dt^\alpha}\right)$ to
 308 a white noise forcing. The general behaviour of responses to linear operators is determined by
 309 their impulse response (Green's) functions $G_{\alpha,h}$ that satisfy:

$$310 \quad \left(\frac{d^{h+\alpha}}{dt^{h+\alpha}} + \frac{d^\alpha}{dt^\alpha}\right)G_{\alpha,h} = \delta(t) \quad (9)$$

311 (Lovejoy 2022)), where $\delta(t)$ is the Dirac ("delta") function. $G_{\alpha,h}$ can be expressed in terms of
 312 "generalized exponentials" or Mittag-Leffler functions $e_{h,h+\alpha}$ as:

$$313 \quad G_{\alpha,h}(t) = \begin{cases} t^{h-1+\alpha} e_{h,h+\alpha}(-t^h) = t^{\alpha-1} \sum_{n=1}^{\infty} (-1)^{n+1} \frac{t^{nh}}{\Gamma(\alpha + nh)}; & t \geq 0 \\ 0; & t < 0 \end{cases} \quad ; \quad \alpha \geq 0; \quad 0 \leq h \leq 2$$

314 (10)

$$315 \quad e_{a,b}(z) = \sum_{n=0}^{\infty} \frac{z^n}{\Gamma(an + b)}$$

316 (Γ is the gamma function). At small t , the leading order term is therefore $G_{\alpha,h}(t) \approx \frac{t^{\alpha-1+h}}{\Gamma(\alpha + h)}$.

317 The large t (asymptotic) expansion is:

$$318 \quad G_{\alpha,h}(t) = t^{\alpha-1} \sum_{n=0}^{\infty} \frac{(-1)^n}{\Gamma(\alpha - nh)} t^{-nh}; \quad t \gg 1 \quad (11)$$

319 (Podlubny 1999). Whereas the small t expansion is $t^{\alpha-1}$ times terms of positive powers of h ,
 320 the large t expansion is in terms of $t^{\alpha-1}$ times terms in negative powers of h , with leading term

321 $G_{\alpha,h}(t) = \frac{t^{\alpha-1}}{\Gamma(\alpha)}$. Unless $h = 0$, $G_{\alpha,h}(t)$ therefore transitions between two different power laws.

322 The special case $h = 0$ that applies to the temperature and turnover forcings (eq. 2), corresponds

323 to the pure power law $G_{\alpha,0}(t) = \frac{t^{\alpha-1}}{\Gamma(\alpha)}$. $G_{\alpha,h}$ has the property that if it is (fractionally) integrated
 324 ζ times, the result is just $G_{\alpha+\zeta,h}$. As explained in appendix A, $G_{\alpha,h}$ is useful for numerical
 325 simulations.

326 At a typical highest resolution of global datasets of 1Myr time scales, for example a
 327 bolide strike (During et al., 2022; Alvarez et al., 1980), supernova or gamma ray burst (Fields
 328 et al., 2020), or even much slower hyperthermal event such as PETM (McInerney & Wing,
 329 2011 ; Gingerich, 2006) or Cenomanian-Turonian Event (Eaton et al., 1997 ; Meyers et al.,
 330 2012 ; Venckutė-Aleksienė et al., 2018) is effectively an impulse, so that $G_{\alpha,h}(t)$ could be
 331 considered as the response to a short time scale stressor such as meteorite or asteroid impact or
 332 extensive volcanic eruption episode. This impulse response property is desirable, since the
 333 global stratigraphic stages and substages are defined based on the episodes of turnover, which
 334 implies that at the measurement scales of million years or more, most of turnover is
 335 intermittent—near instantaneous or impulse-like (Foote, 2005 ; Foote, 1994). Figs. 1, 2 show
 336 the impulse response functions for the empirical parameters estimated in SL ($\alpha \approx 0.25$, $h \approx 0.5$).
 337 Since the equations are linear, these impulse responses will be superposed onto the stochastic
 338 white noise driven responses. We could remark that the power law decay of the impulse
 339 responses is much slower than that of conventionally assumed exponential decays. This means
 340 that our model predicts that there are long term impacts of bolide catastrophic events. This is
 341 in accord – for example - with the findings of (Krug et al., 2009; Krug & Jablonski, 2012), that
 342 the K-Pg mass extinction caused by the effects of Chixulub asteroid impact changed long-term
 343 origination rates and their spatial distribution, that persists today, 66 million years after the
 344 event, in accord with this long memory feature of the FMEM model.

345 2.3. Solving the model

346 Fractional derivatives are generally convolutions (with power laws, eq. 3) and therefore
 347 according to the range of integration of the convolution the fractional derivatives and integrals
 348 will be different. Different convolution ranges therefore correspond to different definitions of
 349 fractional derivatives. Most often (e.g. the Riemann-Liouville and Caputo fractional
 350 derivatives), the latter are taken from time = 0 to t in which case the initial conditions are
 351 important and dealing with them is technically somewhat complex. In these cases, the main
 352 tool is the Laplace transform.

353 Here however, we consider statistically stationary white noise forcing that starts at time
 354 = $-\infty$. In this case, we can use the ‘‘Weyl’’ fractional derivative (a convolution from $-\infty$ to
 355 t , eq. 3) whose Fourier transform (‘‘F.T.’’) is particularly simple:

$$356 \quad \frac{d^h}{dt^h} \overset{F.T.}{\leftrightarrow} (i\omega)^h \quad (12)$$

357 If we Fourier transform (denoted with a tilde), equations 1, 2, we can write the model in
 358 matrix form as:

$$359 \quad \underline{\tilde{S}}(\omega) = (i\omega\tau)^{-\alpha} \underline{\tilde{F}}(\omega) \begin{pmatrix} 1 & 0 \\ 0 & 1 \\ s_T & s_E \end{pmatrix} \begin{pmatrix} \tilde{\gamma}_T \\ \tilde{\gamma}_E \end{pmatrix}$$

$$360 \quad \underline{\tilde{S}}(\omega) = \begin{pmatrix} \tilde{T} \\ \tilde{E}_+ \\ \tilde{D} \end{pmatrix}$$

$$361 \quad \underline{\tilde{F}}(\omega) = \begin{pmatrix} (i\omega\tau)^{-h} & 0 & 0 \\ 0 & 1 & 0 \\ 0 & 0 & \frac{1}{1+(i\omega\tau)^h} \end{pmatrix} \quad (13)$$

362 (the single underline indicates a vector, the double underline, a matrix).

363 As noted above, the D forcing is a linear combination of white noises (eq. 7), so that the
 364 sum on the RHS of eqs. 7, 13 is a correlated white noise. However, from the data (see fig. 4),
 365 we see that E_+ , T are themselves correlated. We therefore rewrite the model in terms of two
 366 statistically independent ($\langle \gamma_1 \gamma_2 \rangle = 0$) unit ($\langle \gamma_1^2 \rangle = \langle \gamma_2^2 \rangle = 1$) white noise drivers γ_1, γ_2 :

$$367 \quad \begin{pmatrix} \underline{\gamma}_T \\ \underline{\gamma}_E \end{pmatrix} = \begin{pmatrix} \sigma_T & 0 \\ 0 & \sigma_E \end{pmatrix} \begin{pmatrix} 1 & 0 \\ \rho_E & \sqrt{1-\rho_E^2} \end{pmatrix} \begin{pmatrix} \gamma_1 \\ \gamma_2 \end{pmatrix} \quad (14)$$

368 So that:

$$369 \quad \sigma_T^2 = \langle \gamma_T^2 \rangle; \quad \sigma_E^2 = \langle \gamma_E^2 \rangle; \quad \rho_E = \frac{\langle \gamma_T \gamma_E \rangle}{\sigma_T \sigma_E}; \quad \langle \gamma_T \rangle = \langle \gamma_E \rangle = 0 \quad (15)$$

370 Where σ_T is the standard deviation of γ_T , σ_E of γ_E and ρ_E is the T, E_+ correlation. Eq. 14 is the
 371 standard Cholesky decomposition of correlated random variables, noises.

372 Fourier transforming eq. 14 and using eq. 13, we can write the model as:

$$373 \quad \underline{\tilde{S}}(\omega) = (i\omega\tau)^{-\alpha} \underline{F}(\omega) \underline{\underline{\sigma}} \underline{\underline{\rho}} \underline{\tilde{\gamma}} \quad (16)$$

$$374 \quad \underline{\underline{\sigma}} = \begin{pmatrix} \sigma_T & 0 & 0 \\ 0 & \sigma_E & 0 \\ 0 & 0 & \sigma_D \end{pmatrix}; \quad \underline{\underline{\rho}} = \begin{pmatrix} 1 & 0 \\ \rho_E & \sqrt{1-\rho_E^2} \\ \rho_D & \text{sgn}(r)\sqrt{1-\rho_D^2} \end{pmatrix}; \quad \underline{\tilde{\gamma}} = \begin{pmatrix} \tilde{\gamma}_1 \\ \tilde{\gamma}_2 \end{pmatrix}$$

375 Where the parameters:

$$376 \quad \sigma_D = s_T \sigma_T \sqrt{1 + 2\rho_E r + r^2}; \quad r = \frac{s_E \sigma_E}{s_T \sigma_T} \quad (17)$$

$$\rho_D = \frac{1 + r\rho_E}{\sqrt{1 + 2r\rho_E + r^2}}$$

377 depend on both the driver statistics (σ_T, σ_E and ρ_E) and the model parameters s_T, s_E . While σ_D
 378 does parametrize the amplitude of the diversity fluctuations, unlike σ_T, σ_E (that must be ≥ 0), it

379 is not a true standard deviation: if $s_T < 0$ it will be negative. Similarly, we will see that ρ_D
 380 determines the D , E_+ and D , T correlations but is not itself a correlation coefficient and it
 381 depends on the sign of the ratio r .

382

383 2.3 Stochastic response to white noise forcing

384 2.3.1 Scaling processes: fGn, fBm

385 We are interested in the statistical properties of the solutions $\tilde{X}(\omega)$. These can be
 386 expressed in terms of fGn, fBm and ffRn (fractionally integrated fractional Relaxation noises)
 387 processes. Before discussing the full statistics that includes the cross correlations, let us
 388 therefore discuss their statistics.

389 Let us start with the scaling processes T , E_+ that are of the form:

$$390 \quad \frac{d^{\alpha_x} X}{dt^{\alpha_x}} = \gamma \stackrel{F.T.}{\leftrightarrow} (i\omega)^{\alpha_x} \tilde{X} = \tilde{\gamma} \quad (18)$$

391 For the statistics, we can determine the power spectrum:

$$392 \quad E_x(\omega) = \left\langle |\tilde{X}|^2 \right\rangle = \frac{1}{2\pi} |\omega|^{-\beta_x}; \quad \beta_x = 2\alpha_x \quad (19)$$

393 Where β_x is the spectral exponent and we have used the fact that the spectrum of a Gaussian
 394 white noise is flat:

$$395 \quad \left\langle |\tilde{\gamma}(\omega)|^2 \right\rangle = \frac{1}{2\pi} \langle \gamma^2 \rangle = \frac{1}{2\pi} \quad (20)$$

396 $E_x(\omega)$ is thus the basic form of the T , E_+ spectra. From the Wiener-Khintchin theorem, the
 397 (real space) autocorrelation function $R_X(\Delta t)$ is the inverse transform:

$$R_X(\Delta t) = \langle X(t)X(t-\Delta t) \rangle \propto \Delta t^{H_X} \stackrel{F.T.}{\leftrightarrow} \tilde{R}_X(\omega) = E_X(\omega) = \langle |X(\omega)|^2 \rangle \propto |\omega|^{-\beta_X}; \quad H_X = \frac{\beta_X - 1}{2} = \frac{\alpha_X}{2}$$

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(21)

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The technical difficulty is that due to a low frequency divergence, the inverse transform of pure power spectra (eq. 19) only converges for $\beta_X < 1$ (i.e. $\alpha_X < 1/2$, $H_X < 0$); this is the fGn regime appropriate for E_+ . Even here, $R_X(\Delta t)$ is infinite for $\Delta t = 0$. Since $R_X(0)$ is the variance, fGn processes are (like the white noise special case $\alpha_X = 0$) generalized functions that must be averaged (integrated) over finite intervals in order to represent physical processes. Averaging to yield a finite resolution process is adequate for $\beta_X > -1$ ($\alpha_X > -1/2$, $H_X > -1$) so that the fGn process is defined for $-1 < \beta_X < 1$ (i.e. $-1/2 < \alpha_X < 1/2$, $-1 < H_X < 0$). After averaged over a finite resolution τ_r : X_{τ_r} with the result $\langle X_{\tau_r}^2 \rangle^{1/2} \propto \tau_r^{H_X}$ and since $H_X < 0$ the data will be highly sensitive to the resolution τ_r .

When $\alpha_X \geq 1/2$, the low frequency divergences imply that the $X(t)$ process is nonstationary (the process generally “wanders off to plus or minus infinity”). However, for $1 < \beta_X < 3$ (i.e. $1/2 < \alpha_X < 3/2$, $0 < H_X < 1$); this is the range appropriate for T : $H_T \approx 0.25$, $\beta_T \approx 3/2$), it’s increments are (stationary) fGn processes, this regime defines the fBm process. Finally, since all physical scaling processes exist over finite ranges of scale, there will be finite outer (longest) time scale (smallest frequency) so that truncating the spectrum at low frequencies (as for the ffRn processes, see below) leads to an overall stationary process.

When analysing paleo series, it is convenient to analyze the statistics in real space, the main reason being that these are easier to interpret (the difficulty in interpretation is the cause of the quadrillion error in climate temperature spectra that was only recently discovered (Lovejoy 2015)). An additional reason is that paleo series are typically not available at uniform sampling / averaging intervals making the spectrum more difficult to estimate.

421 We have already noted that the autocorrelation functions are only adequate for $H_X < 0$
 422 ($\alpha_X < 1/2$, $\beta_X < 1$), this is why when $0 < H_X < 1$, it is conventional to define fluctuations using
 423 differences $\Delta X(\Delta t) = X(t - \Delta t) - X(t)$, which are stationary over this range. Differences avoid
 424 low frequency divergences but will still have high frequency divergences when $H_X < 0$. In order
 425 to avoid the problems at both small scale (resolution dependencies) and at large scales
 426 (nonstationarity), it is convenient to use Haar fluctuations. Over the interval Δt the Haar
 427 fluctuation $\Delta X(\Delta t)$ is defined as the difference between the average of the first and second
 428 halves of the interval.

$$429 \quad \left\langle \Delta X(\Delta t)^2 \right\rangle^{1/2} \propto \Delta t^{H_X} \leftrightarrow E_X(\omega) \propto \omega^{-\beta_X}; \quad \begin{array}{l} -1 < H_X < 1 \\ -1 < \beta_X < 3 \\ \beta_X = 2H_X + 1 \end{array} \quad (22)$$

430 (valid for Haar fluctuations). Over the indicated range of parameters, the Haar fluctuations are
 431 stationary and are independent of the resolution.

432 Comparing eq. 7 and 2 we find:

$$433 \quad \begin{aligned} \left\langle \Delta E_+(\Delta t)^2 \right\rangle^{1/2} &\propto \Delta t^{H_E}; \quad H_E = \alpha - \frac{1}{2} \\ \left\langle \Delta T(\Delta t)^2 \right\rangle^{1/2} &\propto \Delta t^{H_T}; \quad H_T = h + \alpha - \frac{1}{2} \end{aligned} \quad (23)$$

434 *2.3.2 Two scaling regimes: fRn, ffRn*

435

436 From eq. 8, 9, the basic Fourier transforms of ffRn processes and their impulse responses
 437 are:

$$438 \quad \tilde{U}_{\alpha,h}(\omega) = \frac{\tilde{\gamma}}{(i\omega)^\alpha (1+(i\omega)^h)}; \quad \tilde{G}_{\alpha,h}(\omega) = \frac{1}{(i\omega)^\alpha (1+(i\omega)^h)}; \quad 0 < \alpha < 1/2; \quad 0 < h < 2$$

439 (24)

440 The fractional Relaxation noise (fRn) process is the special case where $\alpha = 0$. The ffRn power
 441 spectrum is therefore:

$$442 \quad E_{\alpha,h}(\omega) = \left\langle \left| \tilde{U}_{\alpha,h} \right|^2 \right\rangle = \frac{1}{2\pi |\omega|^{2\alpha} \left| 1 + (i\omega)^h \right|^2} \quad (25)$$

443 $E_{\alpha,h}(\omega)$ is thus the basic form of the D spectrum.

444 The full statistical properties of ffRn processes (including series expansions) are
 445 discussed in (Lovejoy 2022), however for our purposes, the low and high frequency scaling
 446 exponents are sufficient. For these, eq. 25, yields:

$$447 \quad E_{\alpha,h}(\omega) \propto |\omega|^{-\beta}; \quad \begin{array}{ll} \beta_l = 2\alpha; & \omega \ll 1 \\ \beta_h = 2(\alpha + h); & \omega \gg 1 \end{array} \quad (26)$$

448 (“ h ” for high frequency, “ l ” for low frequency). In order to obtain the basic fluctuation
 449 statistics, it is sufficient to apply eq. 22 to each regime separately. Indeed, more generally,
 450 “Tauberian theorems” (e.g. (Feller 1971)) imply that if the spectrum is a power law over a wide
 451 enough range, then the corresponding (second order) real space statistics will also be scaling.
 452 Therefore:

$$453 \quad \left\langle \Delta U_{\alpha,h}(\Delta t)^2 \right\rangle^{1/2} \begin{array}{ll} \propto \Delta t^{H_l}; & H_l = \alpha - \frac{1}{2}; \quad \Delta t \gg 1 \\ \propto \Delta t^{H_h}; & H_h = \alpha + h - \frac{1}{2}; \quad \Delta t \ll 1 \end{array} \quad (27)$$

454 Using the empirical values $\alpha \approx 0.25$, $h \approx 0.5$, we see E_+ is a fractional Gaussian noise and T is
 455 an fBm process. Also, we find (c.f. eqs. 7, 27) that $H_D \approx H_T$ ($\Delta t \ll \tau$) and $H_D \approx H_E$ ($\Delta t \gg \tau$).
 456

457 2.4. The full model statistics: spectra, correlations:

458 *2.4.1 The basic model:*

459 The model is linear and has stationary Gaussian (white noise) forcing (T, E_+), therefore
 460 D, E_-, E, O are also Gaussian so that their statistics are determined by spectra and cross-spectra
 461 – or equivalently in real space (via the Wiener-Khintchin theorem), by the autocorrelations and
 462 cross-correlations:

$$463 \quad R_{ij}(\Delta t) = \langle S_i(t) S_j(t - \Delta t) \rangle \stackrel{F.T.}{\leftrightarrow} \tilde{R}_{ij}(\omega) = \langle \tilde{S}_i(\omega) \tilde{S}_j^*(\omega) \rangle \quad (28)$$

464 (the diagonal terms are the spectra of the components: $\tilde{R}_{ii}(\omega) = E_i(\omega)$). In matrix notation:

$$465 \quad \begin{aligned} \underline{\underline{\tilde{R}}}(\omega) &= \langle \underline{\underline{\tilde{S}}} \underline{\underline{\tilde{S}}}^{T*} \rangle = |\omega\tau|^{-2\alpha} \underline{\underline{F}}(\omega) \underline{\underline{\sigma}} \underline{\underline{\rho}} \langle \underline{\underline{\tilde{\gamma}}} \underline{\underline{\tilde{\gamma}}}^{T*} \rangle \underline{\underline{\rho}}^{T*} \underline{\underline{\sigma}}^{T*} \underline{\underline{F}}(\omega)^{T*} \\ &= \frac{|\omega\tau|^{-2\alpha}}{2\pi} \underline{\underline{F}}(\omega) \underline{\underline{\sigma}} \underline{\underline{\rho}} \underline{\underline{\rho}}^{T*} \underline{\underline{\sigma}}^{T*} \underline{\underline{F}}(\omega)^{T*} \end{aligned} \quad (29)$$

466 Where we have used:

$$467 \quad \langle \underline{\underline{\tilde{\gamma}}} \underline{\underline{\tilde{\gamma}}}^{T*} \rangle = \left\langle \begin{pmatrix} \tilde{\gamma}_1 \\ \tilde{\gamma}_2 \end{pmatrix} \begin{pmatrix} \tilde{\gamma}_1 & \tilde{\gamma}_2 \end{pmatrix} \right\rangle = \frac{1}{2\pi} \begin{pmatrix} 1 & 0 \\ 0 & 1 \end{pmatrix} = \frac{1}{2\pi} \mathbf{1} \quad (30)$$

468 The key correlation matrix (from eq. 16) is:

$$469 \quad \underline{\underline{\rho}} \underline{\underline{\rho}}^{T*} = \begin{pmatrix} 1 & \rho_{TE} & \rho_{TD} \\ \rho_{TE} & 1 & \rho_{ED} \\ \rho_{TD} & \rho_{ED} & 1 \end{pmatrix} \quad (31)$$

470 Where

$$471 \quad \rho_{TE} = \rho_E; \quad \rho_{TD} = \rho_D; \quad \rho_{ED} = \rho_E \rho_D + \text{sgn}(r) \sqrt{1 - \rho_E^2} \sqrt{1 - \rho_D^2} \quad (32)$$

472

473 and

$$474 \quad \underline{\underline{\sigma}} \underline{\underline{\rho}} \underline{\underline{\rho}}^{T*} \underline{\underline{\sigma}}^{T*} = \begin{pmatrix} \sigma_T^2 & \rho_{TE} \sigma_T \sigma_E & \rho_{TD} \sigma_D \sigma_T \\ \rho_{TE} \sigma_T \sigma_E & \sigma_E^2 & \rho_{DE} \sigma_E \sigma_D \\ \rho_{TD} \sigma_D \sigma_T & \rho_{DE} \sigma_E \sigma_D & \sigma_D^2 \end{pmatrix} \quad (33)$$

475

476 *2.4.2 Closing the model: the diagnostic equation for E:*

477 Before writing down the final spectra, let's close the system with the help of the
478 diagnostic equation that allows us to determine E_- from D (and hence E , O , eq. 6).

479 The Fourier transform of the diagnostic equation (eq. 5) is:

$$480 \quad \tilde{E}_- = \left(\frac{\tau_D}{\tau} \right) (i\omega\tau) \tilde{D} \quad (34)$$

481 Therefore the full system is:

$$482 \quad \begin{pmatrix} \tilde{T} \\ \tilde{E}_+ \\ \tilde{D} \\ \tilde{E}_- \end{pmatrix} = (i\omega\tau)^{-\alpha} \begin{pmatrix} (i\omega\tau)^{-h} & 0 & 0 & 0 \\ 0 & 1 & 0 & 0 \\ 0 & 0 & \frac{1}{1+(i\omega\tau)^h} & 0 \\ 0 & 0 & 0 & \frac{i\omega\tau}{1+(i\omega\tau)^h} \end{pmatrix} \begin{pmatrix} \sigma_T & 0 & 0 & 0 \\ 0 & \sigma_{E_+} & 0 & 0 \\ 0 & 0 & \sigma_D & 0 \\ 0 & 0 & 0 & \frac{\tau_D}{\tau} \end{pmatrix} \begin{pmatrix} 1 \\ \rho_E \\ \rho_D \\ \rho_D \end{pmatrix} \begin{pmatrix} 0 \\ \sqrt{1-\rho_E^2} \\ \text{sgn}(r)\sqrt{1-\rho_D^2} \\ \text{sgn}(r)\sqrt{1-\rho_D^2} \end{pmatrix} \begin{pmatrix} \gamma_1 \\ \gamma_2 \end{pmatrix} \quad (35)$$

484 From this we can find E , O :

$$485 \quad \begin{aligned} \tilde{E} &= \frac{1}{2} (\tilde{E}_+ - \tilde{E}_-) \\ \tilde{O} &= \frac{1}{2} (\tilde{E}_+ + \tilde{E}_-) \end{aligned} \quad (36)$$

486 The explicit formulae for E_{\pm} are:

$$487 \quad \begin{aligned} \tilde{E}_+ &= (i\omega\tau)^{-\alpha} \left[\sigma_{E_+} \left(\rho_E \tilde{\gamma}_1 + \sqrt{1-\rho_E^2} \tilde{\gamma}_2 \right) \right] \\ \tilde{E}_- &= (i\omega\tau)^{-\alpha} \frac{i\omega\tau_D}{1+(i\omega\tau)^h} \left(\rho_D \tilde{\gamma}_1 + \text{sgn}(r) \sqrt{1-\rho_D^2} \tilde{\gamma}_2 \right) \end{aligned} \quad (37)$$

488 The overall final statistics are:

$$\begin{aligned}
\tilde{\underline{\underline{R}}}(\omega) &= \begin{pmatrix} \langle \tilde{T}^2 \rangle & \langle \tilde{T} \tilde{E}_+^* \rangle & \langle \tilde{T} \tilde{D}^* \rangle & \langle \tilde{T} \tilde{E}_-^* \rangle \\ \langle \tilde{E}_+ \tilde{T}^* \rangle & \langle \tilde{E}_+^2 \rangle & \langle \tilde{E}_+ \tilde{D}^* \rangle & \langle \tilde{E}_+ \tilde{E}_-^* \rangle \\ \langle \tilde{D} \tilde{T}^* \rangle & \langle \tilde{D} \tilde{E}_+^* \rangle & \langle \tilde{D}^2 \rangle & \langle \tilde{D} \tilde{E}_-^* \rangle \\ \langle \tilde{E}_- \tilde{T}^* \rangle & \langle \tilde{E}_- \tilde{E}_-^* \rangle & \langle \tilde{E}_- \tilde{D}^* \rangle & \langle \tilde{E}_-^2 \rangle \end{pmatrix} \\
&= |\omega\tau|^{-2\alpha} \begin{pmatrix} |\omega\tau|^{-2h} \sigma_T^2 & (i\omega\tau)^{-h} \rho_{TE} \sigma_T \sigma_E & \frac{\rho_{TD} \sigma_D \sigma_T}{(i\omega\tau)^h (1 + (-i\omega\tau)^h)} & \frac{\rho_{TD} \tau_D \sigma_T (-i\omega\tau)^{1-h}}{\tau (1 + (i\omega\tau)^h)} \\ (-i\omega\tau)^{-h} \rho_{TE} \sigma_T \sigma_E & \sigma_E^2 & \frac{\rho_{ED} \sigma_E \sigma_D}{(1 + (-i\omega\tau)^h)} & \frac{\rho_{ED} \tau_D \sigma_E (-i\omega\tau)}{\tau (1 + (i\omega\tau)^h)} \\ \frac{\rho_{TD} \sigma_D \sigma_T}{(-i\omega\tau)^h (1 + (i\omega\tau)^h)} & \frac{\rho_{ED} \sigma_E \sigma_D}{(1 + (i\omega\tau)^h)} & \frac{\sigma_D^2}{|1 + (i\omega\tau)^h|^2} & \frac{\sigma_D \tau_D (-i\omega\tau)}{\tau |1 + (i\omega\tau)^h|^2} \\ \frac{\rho_{TD} \tau_D \sigma_T (i\omega\tau)^{1-h}}{\tau (1 + (i\omega\tau)^h)} & \frac{\rho_{ED} \tau_D \sigma_E (i\omega\tau)}{\tau (1 + (i\omega\tau)^h)} & \frac{\sigma_D \tau_D (i\omega\tau)}{\tau |1 + (i\omega\tau)^h|^2} & \frac{\tau_D^2 |\omega\tau|^2}{\tau^2 |1 + (i\omega\tau)^h|^2} \end{pmatrix}
\end{aligned} \tag{38}$$

489

490

491 Using eqs. 36, 37, the spectra of E , O can be determined:

$$\begin{aligned}
\langle \tilde{E}^2 \rangle &= \frac{1}{4} \left(\langle \tilde{E}_+^2 \rangle + \langle \tilde{E}_-^2 \rangle - 2 \langle \tilde{E}_+ \tilde{E}_-^* \rangle \right) \approx \frac{1}{4} \langle \tilde{E}_+^2 \rangle \\
\langle \tilde{O}^2 \rangle &= \frac{1}{4} \left(\langle \tilde{E}_+^2 \rangle + \langle \tilde{E}_-^2 \rangle + 2 \langle \tilde{E}_+ \tilde{E}_-^* \rangle \right) \approx \frac{1}{4} \langle \tilde{E}_+^2 \rangle
\end{aligned} \tag{39}$$

492

493 The far right approximation can be seen from eq. 37 using the fact that τ_D is the resolution of

494 the series so that for the full range of empirically accessible frequencies, we have $\omega\tau_D < 1$. In

495 addition, since $\tau > \tau_D$, the factor $\left| \omega\tau_D / (1 + (i\omega\tau)^h) \right| \ll 1$.

496 **3. The properties of the model**

497 3.1 Scaling properties

498 *3.1.1 High and low frequency exponents*

499 In order to interpret the statistics (eqs. 38, 39) in real space, it suffices to use the fact that
 500 Fourier scaling implies real space scaling and to use the above relations between real space and
 501 Fourier scaling exponents (eq. 22). In matrix form, the spectral exponents are therefore:

502
$$\beta_h = \begin{pmatrix} 2(\alpha+h) & 2\alpha+h & 2(\alpha+h) & 2(\alpha+h)-1 \\ 2\alpha+h & 2\alpha & 2\alpha+h & 2\alpha+h-1 \\ 2(\alpha+h) & 2\alpha+h & 2(\alpha+h) & 2(\alpha+h)-1 \\ 2(\alpha+h)-1 & 2\alpha+h-1 & 2(\alpha+h)-1 & 2(\alpha+h-1) \end{pmatrix} \quad (40)$$

503
$$\beta_l = \begin{pmatrix} 2(\alpha+h) & 2\alpha+h & 2\alpha+h & 2\alpha+h-1 \\ 2\alpha+h & 2\alpha & 2\alpha & 2\alpha-1 \\ 2\alpha+h & 2\alpha & 2\alpha & 2\alpha-1 \\ 2\alpha+h-1 & 2\alpha-1 & 2\alpha-1 & 2(\alpha-1) \end{pmatrix}$$

504 (The elements correspond to T, E_+, D, E_- left to right, top to bottom). Using the relationship
 505 between H and β (eq. 22), the high and low frequency (here small and large times, t) have
 506 exponents:

507
$$H_h = \begin{pmatrix} \alpha+h-\frac{1}{2} & \alpha+\frac{h-1}{2} & \alpha+h-\frac{1}{2} & \alpha+h-1 \\ \alpha+\frac{h-1}{2} & \alpha-\frac{1}{2} & \alpha+\frac{h-1}{2} & \alpha+\frac{h}{2}-1 \\ \alpha+h-\frac{1}{2} & \alpha+\frac{h-1}{2} & \alpha+h-\frac{1}{2} & \alpha+h-1 \\ \alpha+h-1 & \alpha+\frac{h}{2}-1 & \alpha+h-1 & \alpha+h-\frac{3}{2} \end{pmatrix} \quad (41)$$

508 While at low frequencies large Δt (i.e. large lags) we have:

509

510
$$H_l = \begin{pmatrix} \alpha + h - \frac{1}{2} & \alpha + \frac{h-1}{2} & \alpha + \frac{h-1}{2} & \alpha + \frac{h}{2} - 1 \\ \alpha + \frac{h-1}{2} & \alpha - \frac{1}{2} & \alpha - \frac{1}{2} & \alpha - 1 \\ \alpha + \frac{h-1}{2} & \alpha - \frac{1}{2} & \alpha - \frac{1}{2} & \alpha - 1 \\ \alpha + \frac{h}{2} - 1 & \alpha - 1 & \alpha - 1 & \alpha - \frac{3}{2} \end{pmatrix} \quad (42)$$

511

512 We should add here that since E , O are linear combinations of E_+ , E_- , their exponents will be the
513 maximum of those of E_+ , E_- , so that:

514
$$H_{h,E_{\pm}} = \max\left(\alpha - \frac{1}{2}, \alpha + h - \frac{3}{2}\right) = \alpha - \frac{1}{2}; \quad h < 1 \quad (43)$$

$$H_{l,E_{\pm}} = \max\left(\alpha - \frac{1}{2}, \alpha - \frac{3}{2}\right) = \alpha - \frac{1}{2}$$

515 We see that for the physically relevant parameters, $H = \alpha - 1/2 = -0.25$ for both E , O , over the
516 whole range (close to the data, see SL and fig. 3).

517 To get a concrete idea of the implications of model, let's use the rough empirical
518 estimates from SL of $\alpha = 0.25$, $h = 0.5$. Plugging these values into eqs. 41, 42, we obtain:

519
$$H_h = \begin{pmatrix} 0.25 & 0 & 0.25 & -0.25 \\ 0 & -0.25 & 0 & -0.5 \\ 0.25 & 0 & 0.25 & -0.25 \\ -0.25 & -0.5 & -0.25 & -0.75 \end{pmatrix} \quad H_l = \begin{pmatrix} 0.25 & 0 & 0 & -0.5 \\ 0 & -0.25 & -0.25 & -0.75 \\ 0 & -0.25 & -0.25 & -0.75 \\ -0.5 & -0.75 & -0.75 & -1.25 \end{pmatrix} \quad (44)$$

520 (again, for T , E_+ , D , E_- left to right, top to bottom). We can see that the Haar fluctuations will
521 be useful for all the series over the whole range of frequencies/scales, the only exception being
522 $\Delta D(\Delta t)$ at long lags ($H_l < -1$, lower right corner of the H_l matrix with $H_l < -1$). In this case, the
523 Haar fluctuations “saturate” and the spurious (limiting) value $H_l = -1$ is obtained.

524

525 *3.1.2 Normalized Correlations*

526 The cross spectra and cross covariances (eq. 38) can be used to determine the normalized
527 correlations that were estimated in SL:

$$528 \quad \rho_{jk}(\Delta t) = \frac{\langle \Delta S_j(\Delta t) \Delta S_k(\Delta t) \rangle}{\langle \Delta S_j(\Delta t)^2 \rangle^{1/2} \langle \Delta S_k(\Delta t)^2 \rangle^{1/2}} \quad (45)$$

529
530 (Haar fluctuations). However, from eqs. 41, 42, we find that their exponents (whether at high
531 or low frequencies) are $2H_{jk} - (H_{jj} + H_{kk}) = 0$ i.e. they are *not* power laws and only vary at
532 sub power law rates, they are therefore nontrivial (i.e. they are significant) over the whole range
533 of Δt . Since there are six series (T, E, D, O, E_+, E_-) there are 15 pairs whose fluctuation
534 correlations may be determined over the observed range of $3 \approx < \Delta t \approx < 400$ Myrs, see fig. 4.
535 The key correlations are those that correspond to the model parameters: $\rho_E = \rho_{TE}$, $\rho_D = \rho_{TD}$,
536 see below. We can already see that the correlations are quite noisy, a consequence of the low
537 resolution and variable sampling of the series. In order to make a proper model - data
538 comparison, we therefore turn to numerical simulations.

539 **4. Numerical simulations:**

540 *4.1 The statistics of the simulated series*

541 The model has two fundamental exponents (α, h), two basic correlations ($\rho_E = \rho_{TE+}$,
542 $\rho_D = \rho_{TD}$) and a cross over time scale τ . The third correlation ρ_{DE} is a derived parameter (eq.
543 32). In addition, there are two amplitude factors σ_T, σ_E but these will depend on the
544 nondimensionalization/normalization of the series; on log-log plots they correspond to an up-
545 down shift and on (normalized) correlation plots, the normalization eliminates them, they will
546 not be considered further.

547 We used the results of SL to fix the values $\alpha = -0.25$, $h = 0.5$, $\tau = 32\text{Myrs}$ (this is the
 548 nearest power of 2 to the slightly larger – but only roughly estimated - value $\tau = 40\text{Myrs}$ in
 549 SL). This leaves the only unknown parameters as the TE and TD correlations ($\rho_E = \rho_{TE+}$,
 550 $\rho_D = \rho_{TD}$), fig. 4.

551 Before comparing the model directly to the (noisy) data we first check that we are able
 552 to numerically reproduce the theoretically expected behaviour. The basic modelling technique
 553 is to use convolutions with various (impulse response) Green’s functions, this is detailed in
 554 appendix A, but follows the methods described in (Lovejoy 2022). The main numerical
 555 problems are the small scales that have singular power law filters that are not trivial to
 556 discretize, and there are some (easier to handle) long time (low frequency) issues.

557 Rather than attempting to rigorously determine optimum parameters, as indicated above,
 558 we fixed the exponents $\alpha = 0.25$, $h = 0.5$ and the crossover scale $\tau = 32\text{Myrs}$. With guidance of
 559 the fig. 4 correlations for ρ_{TE+} , ρ_{TD} and some numerical experimentation, we took $\rho_E = 0.5$,
 560 $\rho_D = -0.1$ (hence $\rho_{TE+} = 0.5$, $\rho_{TD} = -0.1$, $\rho_{DE+} = -0.9$ i.e. the sign of r was taken as negative,
 561 eq. 32). We then performed simulations at a resolution of 250kyrs resolution, with simulation
 562 length of 4 Gyrs ($2^{14} = 16384$ points), shown in fig. 5. We postpone a discussion of the
 563 significance of the correlations to section 4.2.

564 According to the model (see the diagonal elements in eq. 44), the only series with
 565 positive low frequency scaling exponent ($H_l > 0$) is the temperature ($H_l = 0.25$), it indeed shows
 566 “wandering” behaviour (second from the bottom in fig. 5); from the figure, one can see its long
 567 range correlations as low frequency undulations. This is also true for D , but only up to the
 568 cross-over scale ($\approx 32\text{Myrs}$) after which consecutive 32 Myr intervals tend to cancel ($H_l < 0$, eq.
 569 44). The other series are on the contrary “cancelling” ($H_l < 0$, $H_h < 0$) especially E . (eq. 44).
 570 We can also visually make out some of the correlations, but this is clearer at lower resolution
 571 discussed later.

572 On these simulations, we can check that the theoretical scaling is obeyed, this was done
 573 using Haar fluctuations, see fig. 6 where the theory slopes (from eq. 43, 44) are shown as
 574 reference lines. Note that since the Haar analysis “saturates” at $H = -1$, the low frequency $H_l =$
 575 -1.25 value for E . (eq. 44, lower right hand diagonal element) yields a slope -1 (not -1.25), the
 576 other slopes are however accurately estimated. Note that the theory / simulation agreement is
 577 not perfect, mostly because the theory is for the average statistics over an infinite ensemble,
 578 whereas fig. 6 is from a single - albeit large - simulation.

579 We can also work out the 15 correlations as functions of lag, fig. 7. The figure shows
 580 the model parameters $\rho_{TE+} = 0.5$ ($=\rho_E = 0.5$), $\rho_{TD} = -0.1$ ($=\rho_D = -0.1$) as solid black reference
 581 lines and the derived correlation $\rho_{DE+} = -0.9$ (eq. 32) as a dashed reference lines. Also shown
 582 are dashed theory lines for the TE , TO correlations (predicted to be equal to equal to TE_+ at
 583 long lags – eq. 39) and the DE , DO correlations (predicted to be equal to DE_+ , at long lags, see
 584 eq. 39). We can see that the correlations approach the theoretical correlations at large lags,
 585 although the results are somewhat noisy.

586

587 4.2 The statistics of the simulated series resampled at the data sampling times

588 Before making more effort at parameter fitting and comparing the model to data, it is
 589 important to take into account the small number of empirical data points and their irregular
 590 sampling. Fig. 8 shows the result for a simulation with the same parameters, but with a 1 Myr
 591 temporal resolution (right hand side), resampled at the same times as the data (left hand side).
 592 Since the model and data are only expected to have similar statistics, the detailed “bumps” and
 593 “wiggles” are unimportant, but one can nevertheless make out realistic looking variability
 594 including correlations between the series. Note that the model respects causality so that when
 595 there is a large extinction event, that is asymmetric with a rapid upturn being followed by a

596 slower downturn (however, we have followed convention so that the present is at the left and
 597 the past at the right).

598 We can now consider the fluctuation scaling and correlation statistics on the resampled
 599 series and compare them to the both the data and to the results from the same simulations but
 600 at a regular 1 Myr resolution (fig. 9). The figure shows a log-log plot of the RMS fluctuations
 601 as a function of the lag. In order to make the comparison, they were normalized by their
 602 standard deviations, but this is somewhat arbitrary so that the up-down displacement
 603 (corresponding to a different nondimensionalization/normalization) is unimportant. To judge
 604 the realism of the model, the appropriate comparison is between the shapes of the resampled
 605 model output (red) and the data (black). We can see that the two are fairly close although both
 606 model and data are noisy due to the small number of points and the irregular sampling. The
 607 agreement must be assessed not only by allowing for (relative) vertical shifts, but also noting
 608 that the scales on the top D , T comparisons are such that the fluctuations vary only over a small
 609 factors (for the data, factors of ≈ 1.7 for D and ≈ 2 for T) for lags varying over range of about
 610 a factor 100. In comparison, the E_+ , E_- , E , O ranges are closer to factors of 10. Aside from
 611 this, these basic fluctuation statistics are fairly close to the data.

612 The figure also gives important information about the effect of the sampling: compare
 613 the resampled (red) and uniformly sampled analyses (brown). The resampling is particularly
 614 important for E_+ , E_- , E , O although the effects are mostly at small lags for E_+ , E , O but for large
 615 lags for E_- . This information should prove useful in interpreting a variety of real world
 616 extinction and origination data.

617 Finally, we can compare the 15 pairwise correlations (fig. 10). Again, to judge the
 618 realism of the model, compare the red and black correlations. Although – as expected – these
 619 are fairly noisy, we see that the agreement is quite good, significantly, it is generally much
 620 better than the agreement between the uniformly sampled correlations (brown curves) and data

621 (black). By comparing the red (resampled) and brown (uniformly sampled) correlations, we
 622 see that the resampling is especially important for the DE_+ , DO , DE , E_+E_- , $E-O$, OE ,
 623 correlations and to a lesser extent the OE , TE_+ comparisons, for the others it is about the same.
 624 We could note the successful prediction that the E_+E , E_+O , OE correlations that should be ≈ 1
 625 and the $E-E$ correlations that should be ≈ -1 . Interestingly, the prediction that the $E-O$
 626 correlations should be ≈ -1 (eq. 39) is verified with the uniform sampling (i.e. it is indeed a
 627 property of the model), yet, the resampling (red in the lower left graph in fig. 10) makes it >0
 628 and aligns it closely with the observations. In other words, when the pure model predictions
 629 are poor (brown versus black), there are many instances where the effects of nonuniform
 630 sampling are particularly strong so that overall the model explains the data fairly well: overall
 631 6 fluctuation plots (fig. 9) and 15 correlations (fig. 10) with 5 adjustable parameters (α , h , τ ,
 632 ρ_E , ρ_T).

633 4.3 Discussion of the model and physical significance of the correlations

634 The model was motivated by an attempt to model the diversity process as a scaling
 635 cross-over phenomenon with wandering climate (paleo temperature) and stabilizing life
 636 (turnover) scaling drivers. In the course of the model development, it became clear both
 637 theoretically (due to the definition of the diversity, eq. 5) and empirically, that rather than E , O
 638 being fundamental, it rather the turnover E_+ that is fundamental (indeed, the E_+ and E_- statistics
 639 are quite different (figs. 3, 9) and the E_+E_- correlations are nearly zero (figs. 4, 10). In any
 640 event, the model predicted that E , O would follow the E_+ statistics (eq. 39, fig. 3, 9 and the E_+E
 641 and E_+O correlations in fig. 3, 10).

642 A more counterintuitive finding concerns the correlations. To start with, the model
 643 specifies that the diversity is primarily driven by the temperature up until the cross-over scale,
 644 yet the temperature and diversity are negatively correlated over the entire range! Although at
 645 any given time lag, the DT correlation is small (-0.1), it means that there is a (weak) tendency

646 for the diversity fluctuations to decrease when temperature fluctuations increase and visa versa,
 647 but this is not enough to offset the overall temperature control of the diversity that implies that
 648 consecutive temperature fluctuations tend to add up ($H_T = 0.25 > 0$) and this is a stronger overall
 649 effect.

650 There is an additional more subtle effect. Consider that at each scale, the imposed TE_+
 651 correlation is moderate and positive ($\rho_{TE_+} = 0.5$) and together, ρ_{TD} and ρ_{TE_+} (with $r < 0$, eq. 32)
 652 they imply that at each lag, DE_+ is negatively correlated (reaching the theory value $\rho_{DE_+} \approx -0.9$,
 653 at long lags, see the DE_+ correlation, the brown curve in fig. 10). Since the turnover E_+ also
 654 drives the diversity, (eq. 1), at each scale, we thus have a tendency for T and E_+ fluctuations to
 655 increase (or decrease) together but D and E_+ (and hence T and D to have opposite tendencies).
 656 The overall result is that the weak anticorrelation of D with T and D with E_+ at any fixed scale
 657 is still dominated by the stronger effect of T fluctuations growing with scale and dominating
 658 the E_+ driver at lags $< \tau$.

659 We could remark that $\rho_{TE_+} = 0.5 > 0$ indicates a tendency for temperature changes to
 660 “stimulate” the turnover: periods of increasing temperatures tending to be associated with
 661 increasing turnovers and decreasing temperatures with decreasing turnovers. Also there is a
 662 strong anticorrelation between D and E_+ ($\rho_{DE_+} \approx -0.9$, although it seems to nearly disappear
 663 after the nonuniform sampling, see fig. 10, second in the top row) that indicates that increased
 664 turnover decreases with diversity. However over the range of scales that E_+ dominates
 665 dynamics of D (i.e. $\Delta t > \tau$), since $H_{E_+} \approx -0.25 < 0$), successive E_+ fluctuations tend to cancel and
 666 on long time scales, the latter effect is dominant so that $H_D = H_{E_+} \approx -0.25$ – this is a scaling
 667 region of biotic self-regulation.

668 **5. Conclusions:**

669 The driver of macroevolutionary biodiversity has famously been reduced to a dichotomy
 670 between “life” and the “environment”: the “Red Queen” versus “Court Jester” metaphor (Van
 671 Valen 1973); (Barnosky, 2001). Using genus level time series from the Paleobiology Database
 672 (Spiridonov and Lovejoy 2022) (SL) systematically analysed fluctuations in extinction (E),
 673 origination (O) rates, biodiversity (D) and paleo temperatures (T) over the Phanerozoic. They
 674 did this as a function of time scale from the shortest (≈ 3 Myrs) to longest lags available
 675 (≈ 400 Myrs) and their analysis included the correlations of the fluctuations at each scale. They
 676 concluded that T , E , O – the basic climate and life parameters - showed evidence of wide range
 677 scaling, supporting the hypothesis that over this range, there is a single biogeological
 678 “megacclimate” (Lovejoy 2015) regime with no fundamental time scale. However, they found
 679 that D followed the T fluctuations up until a critical time $\tau \approx 40$ Myrs, whereas at longer time
 680 scales, it followed life (E , O): D was a scaling cross-over phenomenon. At the shorter time
 681 scales $\Delta t < \tau$, - like the temperature – the D scaling exponent $H_D \approx +0.25$ (i.e. >0) indicating
 682 that fluctuations tended to grow with scale, leading to “wandering” behaviour. In contrast for
 683 time lags $\Delta t > \tau$, - like E , O , its scaling exponent was $H_D \approx -0.25$ i.e. <0), hence successive
 684 fluctuations tended to cancel, resulting in long time stabilization of diversity by life.

685 In order to clarify our ideas, to better understand the geobiodynamics and to better
 686 understand and quantify the limitations, biases and other data issues, we proposed the simple
 687 model Fractional Macro Evolution Model (FMEM) to reproduce the observations. It is a
 688 model of macroevolutionary biodiversity driven by paleotemperature (the climate proxy) and
 689 the turnover rate ($E_+ = O+E$), the “life” proxy. In order to fit with basic empirical scaling
 690 statistics and theoretical ideas about the macroclimate regime (from time scales of roughly
 691 1Myr to at least 500Myrs), these drivers were taken to be scaling with climate dominating at
 692 short time scales and life at long time scales. Therefore, FMEM suggests a possible way to

693 combine into single stochastic framework both: i) the destabilizing geophysical processes (and
694 possibly astrophysical ((Raup, 1991; Raup, 1992b; Melott & Bambach, 2014 ; Fields et al.,
695 2020)) with ii) the stabilizing, density dependent and self-regulating, biotic processes. The
696 model is specified by a simple parametrization based on two scaling exponents and two
697 pairwise correlations (between T and E_+ and between T and D).

698 The model had two unusual characteristics: first, it was stochastic so that the crossover
699 from climate to life dominance was thus a scaling (power law) not standard exponential (i.e.
700 Markov process type) transition. Stochastic models involve infinite dimensional probability
701 spaces, they are therefore natural model types in systems with huge numbers of degrees of
702 freedom. We believe that they are intrinsically more realistic than strongly nonlinear but
703 deterministic chaos type models (including those that are deterministic but are perturbed by
704 noises). When the intermittency is strong scaling stochastic models must be nonlinear (e.g.
705 multifractal cascade processes), and this can easily be included in further model improvements
706 – the Gaussian forcing (γ_1, γ_2 , eq. 14) need only be replaced by a multifractal one. Here,
707 intermittency was neglected and linear stochastic equations with Gaussian white noise forcings
708 were used (linear stochastic models can often be used even when the underlying dynamics are
709 strongly nonlinear).

710 The other unusual FMEM characteristic was that it a system of fractional differential
711 equations. Unlike the familiar integer ordered differential equations that typically have
712 exponential impulse response functions (Green's functions), fractional equations typically have
713 power law response functions and are natural ways to model scaling processes. These impulse
714 response functions are physical models of bolide impacts and similar nearly instantaneous
715 processes, and we discussed some implications.

716 The model was also highly parsimonious with two scaling exponents and a cross-over
717 time τ determined by the Paleobiology Database data as analyzed SL. These determined the

718 basic scaling characteristics of the 6 series: T , E_+ , D , E . ($= O - E$), O , E . In addition, the model
719 had two correlations that were specified: those between T and E_+ and between T and D . From
720 these, the other 13 pairwise correlations (out the 15 possible pairs of the six series), were
721 implicitly determined and were compared to the data.

722 The fractional derivatives were of the Weyl type so that their Fourier transforms were
723 simple power laws. Since the system was ultimately forced by two Gaussian white noises, only
724 the second order statistics (i.e. the spectra and correlation functions) were needed and these
725 were easily obtained: the basic solutions were fractionally integrated fractional relaxation
726 noises (ffRn) that were recently introduced (Lovejoy 2022). In future, more realistic
727 intermittent (multifractal) forcings could be used instead of the Gaussian white noise. Beyond
728 exhibiting the full solution to the equations with a full statistical characterization, we then
729 implemented the model numerically first verifying the model against the theoretically predicted
730 behaviour. By producing simulations at 1Myr resolution, were able to resample the output at
731 the same irregular sampling times as the biodata base. The statistical characteristics of the
732 results (the 6 scaling curves showing the fluctuations as functions of time scale), plus the 15
733 pairwise correlations as functions of time scale, were all quite close to the data and in several
734 cases, the agreement could be clearly attributed to the limitations, biases, etc. of the data. In
735 particular, this was the case of the DE_+ , DO , DE , E_+E , $E.O$, OE correlations that were much
736 closer to the data following the irregular sampling than with the original model outputs
737 uniformly sampled at 1Myr resolution.

738 Given the model's simplicity, it thus was remarkably realistic. This was fortunate since
739 until higher resolution (global scale) time series become available (e.g. (Fan et al. 2020)), more
740 complex models may not be warranted. In any case, the model was able to help explain some
741 subtle points about the interaction of different correlated series that were also strongly self-

742 correlated over wide ranges of time scales and this with quantitatively and qualitatively
 743 different scaling behaviours (“wandering” versus “cancelling”/self-stabilizing).

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748 **Competing Interests**

749 The authors declare none.

750 **References**

751 Alroy J (2010a) Geographical, environmental and intrinsic biotic controls on
 752 Phanerozoic marine diversification. *Palaeontology*, **53**, 1211-1235.

753 Alroy J (2010b) The shifting balance of diversity among major marine animal
 754 groups. *Science*, **329**, 1191-1194.

755 Alroy J (2015) A more precise speciation and extinction rate estimator.
 756 *Paleobiology*, **41**, 633-639.

757 Alroy J, Aberhan M, Bottjer DJ, Foote M, Fürsich FT, Harries PJ, Hendy AJ, Holland
 758 SM, Ivany LC, Kiessling W (2008) Phanerozoic trends in the global diversity of marine
 759 invertebrates. *Science*, **321**, 97-100.

760 Alroy J, Marshall C, Bambach R, Bezusko K, Foote M, Fürsich F, Hansen TA, Holland
 761 S, Ivany L, Jablonski D, Jacobs DK, Jones DC, Kosnik MA, Lidgard S, Low S, Miller AI,
 762 Novack-Gottshall PM, Olszewski TD, Patzkowsky ME, Raup DM, Roy K, Sepkoski JJJ,
 763 Sommers MG, Wagner PJ, Webber A (2001) Effects of sampling standardization on

764 estimates of Phanerozoic marine diversification. *Proceedings of the National Academy of*
 765 *Sciences*, **98**, 6261-6266.

766 Alvarez LW, Alvarez W, Asaro F, Michel HV (1980) Extraterrestrial cause for the
 767 Cretaceous-Tertiary extinction. *Science*, **208**, 1095-1108.

768 Barnosky AD (2001) Distinguishing the effects of the Red Queen and Court Jester on
 769 Miocene mammal evolution in the northern Rocky Mountains. *Journal of Vertebrate*
 770 *Paleontology*, **21**, 172-185.

771 Bartoszek K, Glémin S, Kaj I, Lascoux M (2017) Using the Ornstein–Uhlenbeck
 772 process to model the evolution of interacting populations. *Journal of Theoretical Biology*,
 773 **429**, 35-45.

774 Beaufort L, Bolton CT, Sarr A-C, Suchéras-Marx B, Rosenthal Y, Donnadiou Y,
 775 Barbarin N, Bova S, Cornuault P, Gally Y (2022) Cyclic evolution of phytoplankton forced
 776 by changes in tropical seasonality. *Nature*, **601**, 79-84.

777 Benton MJ (1995) Diversification and extinction in the history of life. *Science*, **268**,
 778 52-58.

779 Benton MJ (2009) The Red Queen and the Court Jester: species diversity and the
 780 role of biotic and abiotic factors through time. *Science*, **323**, 728-732.

781 Boyle J, T., Sheets H, David,, Wu S-Y, Goldman D, Melchin M, J., Cooper R, A., Sadler
 782 P, M., Mitchell C, E (2013) A re-examination of the contributions of biofacies and
 783 geographic range to extinction risk in Ordovician graptolites. *GFF*, **136**, 38-41.

784 Brayard A, Escarguel G, Bucher H, Monnet C, Brühwiler T, Goudemand N, Galfetti T,
 785 Guex J (2009) Good genes and good luck: ammonoid diversity and the end-Permian mass
 786 extinction. *Science*, **325**, 1118-1121.

787 Caraballoa, T., R. Colucci, and X. Han. 2016. Non-autonomous dynamics of a semi-
 788 Kolmogorov population model with periodic forcing. *Nonlinear Anal. Real.* 31: 661–680.

- 789 Carrillo JD, Faurby S, Silvestro D, Zizka A, Jaramillo C, Bacon CD, Antonelli A (2020)
 790 Disproportionate extinction of South American mammals drove the asymmetry of the
 791 Great American Biotic Interchange. *Proceedings of the National Academy of Sciences*, **117**,
 792 26281-26287.
- 793 Casey MM, Saupe EE, Lieberman BS (2021) The effects of geographic range size and
 794 abundance on extinction during a time of “sluggish” evolution. *Paleobiology*, **47**, 54-67.
- 795 Cornette JL, Lieberman BS (2004) Random walks in the history of life. *Proceedings*
 796 *of the National Academy of Sciences*, **101**, 187-191.
- 797 Crampton JS, Cooper RA, Sadler PM, Foote M (2016) Greenhouse– icehouse
 798 transition in the Late Ordovician marks a step change in extinction regime in the marine
 799 plankton. *Proceedings of the National Academy of Sciences*, **113**, 1498-1503.
- 800 Crampton JS, Meyers SR, Cooper RA, Sadler PM, Foote M, Harte D (2018) Pacing of
 801 Paleozoic macroevolutionary rates by Milankovitch grand cycles. *Proceedings of the*
 802 *National Academy of Sciences*, **115**, 5686-5691.
- 803 Cuthill JFH, Guttenberg N, Budd GE (2020) Impacts of speciation and extinction
 804 measured by an evolutionary decay clock. *Nature*, **588**, 636-641.
- 805 Del Rio Amador, L., and S. Lovejoy. 2021. Using regional scaling for temperature
 806 forecasts with the Stochastic Seasonal to Interannual Prediction System (StocSIPS). *Clim.*
 807 *Dyn.*
- 808 Dowsett HJ, Robinson MM (1998) Application of the modern analog technique
 809 (MAT) of sea surface temperature estimation to middle Pliocene North Pacific planktonic
 810 foraminifer assemblages. *Palaeontologia Electronica*, **1**, 22.
- 811 During MA, Smit J, Voeten DFaE, Berruyer C, Tafforeau P, Sanchez S, Stein KH,
 812 Verdegaal-Warmerdam SJ, Van Der Lubbe JH (2022) The Mesozoic terminated in boreal
 813 spring. *Nature*, **603**, 91-94.

814 Eaton JG, Kirkland JI, Hutchison JH, Denton R, O'Neill RC, Parrish JM (1997)
 815 Nonmarine extinction across the Cenomanian-Turonian boundary, southwestern Utah,
 816 with a comparison to the Cretaceous-Tertiary extinction event. *Geological Society of*
 817 *America Bulletin*, **109**, 560-567.

818 Eldredge N (1985) *Unfinished synthesis: biological hierarchies and modern*
 819 *evolutionary thought*, Oxford University Press, New York.

820 Eldredge N (1989) *Macroevolutionary dynamics: Species, Niches & Adaptive Peaks*,
 821 McGraw-Hill Publishing Company, New York.

822 Eldredge N (2003) The sloshing bucket: how the physical realm controls evolution.
 823 *Evolutionary dynamics: exploring the interplay of selection, accident, neutrality, and*
 824 *function (SFI Studies in the Sciences of Complexity Series)*. Oxford University Press, New
 825 York, 3-32.

826 Erwin DH (2011) Evolutionary uniformitarianism. *Developmental biology*, **357**, 27-
 827 34.

828 Erwin DH (2012) Novelty that change carrying capacity. *Journal of Experimental*
 829 *Zoology Part B: Molecular and Developmental Evolution*, **318**, 460-465.

830 Erwin DH (2016) Wonderful life revisited: chance and contingency in the
 831 Ediacaran-Cambrian radiation. *Chance in evolution*, 279-298.

832 Fan, J.-x., S.-z. Shen, D. H. Erwin, P. M. Sadler, N. MacLeod, Q.-m. Cheng, X.-d. Hou, J.
 833 Yang, X.-d. Wang, Y. Wang, H. Zhang, X. Chen, G.-x. Li, Y.-c. Zhang, Y.-k. Shi, D.-x. Yuan, Q.
 834 Chen, L.-n. Zhang, C. Li, and Y.-y. Zhao. 2020. A high-resolution summary of Cambrian to
 835 Early Triassic marine invertebrate biodiversity. *Science* 367,:272-277.

836 Feller, W. 1971. *An Introduction to probability theory and its applications*, vol. 2.
 837 Wiley, New York.

- 838 Fields BD, Melott AL, Ellis J, Ertel AF, Fry BJ, Lieberman BS, Liu Z, Miller JA, Thomas
839 BC (2020) Supernova triggers for end-Devonian extinctions. *Proceedings of the National*
840 *Academy of Sciences*, **117**, 21008-21010.
- 841 Finnegan S, Payne JL, Wang SC (2008) The Red Queen revisited: reevaluating the
842 age selectivity of Phanerozoic marine genus extinctions. *Paleobiology*, **34**, 318-341.
- 843 Foote M (1994) Temporal variation in extinction risk and temporal scaling of
844 extinction metrics. *Paleobiology*, **20**, 424-444.
- 845 Foote M (2005) Pulsed origination and extinction in the marine realm. *Paleobiology*,
846 **40**, 6-20.
- 847 Gingerich PD (1993) Quantification and comparison of evolutionary rates. *American*
848 *Journal of Science*, **293**, 453-478.
- 849 Gingerich PD (2001) Rates of evolution on the time scale of the evolutionary
850 process. *Microevolution rate, pattern, process*, 127-144.
- 851 Gingerich PD (2006) Environment and evolution through the Paleocene–Eocene
852 thermal maximum. *Trends in ecology & evolution*, **21**, 246-253.
- 853 Gingerich PD (2009) Rates of evolution. *Annual Review of Ecology, Evolution, and*
854 *Systematics*, **40**, 657-675.
- 855 Gould SJ (1990) *Wonderful life: the Burgess Shale and the nature of history*, WW
856 Norton & Company.
- 857 Gould SJ (2001) Contingency. *Palaeobiology II*, 195-198.
- 858 Gould SJ (2002) *The structure of evolutionary theory*, Harvard University Press,
859 Cambridge, MA.
- 860 Gould SJ, Raup DM, Sepkoski Jr JJ, Schopf TJ, Simberloff DS (1977) The shape of
861 evolution: a comparison of real and random clades. *Paleobiology*, 23-40.

- 862 Green W (2006) Loosening the CLAMP: an exploratory graphical approach to the
863 Climate Leaf Analysis Multivariate Program. *Palaeontologia Electronica*, **9**, 1-17.
- 864 Grossman EL, Joachimski MM (2022) Ocean temperatures through the Phanerozoic
865 reassessed. *Scientific Reports*, **12**, 1-13.
- 866 Halliday TJD, Holroyd PA, Gheerbrant E, Prasad GVR, Scanferla A, Beck RMD, Krause
867 DW, Goswami A (2020) Leaving Gondwana: the changing position of the Indian
868 subcontinent in the Global Faunal Network. In: *Biological Consequences of Plate Tectonics*.
869 Springer, pp. 227-249.
- 870 Hannisdal B, Peters SE (2011) Phanerozoic Earth system evolution and marine
871 biodiversity. *Science*, **334**, 1121-1124.
- 872 Harmon LJ, Pennell MW, Henao-Diaz LF, Rolland J, Siple BN, Uyeda JC (2021)
873 Causes and consequences of apparent timescaling across all estimated evolutionary rates.
874 *Annual Review of Ecology, Evolution, and Systematics*, **52**, 587-609.
- 875 Hoffman A (1987) Neutral model of taxonomic diversification in the Phanerozoic: a
876 methodological discussion. In: *Neutral Models in Biology*. Oxford Univ. Press New York,
877 pp. 133-146.
- 878 Huisman, J., and F. J. Weissing. 1999. Biodiversity of plankton by species oscillations
879 and chaos. *Nature* 402(6760):407-410.
- 880 Jablonski D (1986) Background and mass extinctions: the alternation of
881 macroevolutionary regimes. *Science*, **231**, 129-133.
- 882 Jablonski D (2008) Biotic interactions and macroevolution: extensions and
883 mismatches across scales and levels. *Evolution*, **62**, 715-739.
- 884 Jablonski D, Roy K, Valentine JW (2006) Out of the tropics: evolutionary dynamics
885 of the latitudinal diversity gradient. *Science*, **314**, 102-106.

- 886 Jernvall J, Fortelius M (2002) Common mammals drive the evolutionary increase of
887 hypsodonty in the Neogene. *Nature*, **417**, 538-540.
- 888 Khabbazian M, Kriebel R, Rohe K, Ane C (2016) Fast and accurate detection of
889 evolutionary shifts in Ornstein–Uhlenbeck models. *Methods in Ecology and Evolution*, **7**,
890 811-824.
- 891 Kiessling W, Simpson C, Foote M (2010) Reefs as cradles of evolution and sources
892 of biodiversity in the Phanerozoic. *Science*, **327**, 196-198.
- 893 Kocsis AT, Reddin CJ, Alroy J, Kiessling W (2019) The R package divDyn for
894 quantifying diversity dynamics using fossil sampling data. *Methods in Ecology and*
895 *Evolution*, **10**, 735-743.
- 896 Krug AZ, Jablonski D (2012) Long-term origination rates are reset only at mass
897 extinctions. *Geology*, **40**, 731-734.
- 898 Krug AZ, Jablonski D, Valentine JW (2009) Signature of the end-Cretaceous mass
899 extinction in the modern biota. *Science*, **323**, 767-771.
- 900 Lidgard S, Di Martino E, Zágoršek K, Liow LH (2021) When fossil clades ‘compete’:
901 local dominance, global diversification dynamics and causation. *Proceedings of the Royal*
902 *Society B*, **288**, 20211632.
- 903 Lieberman BS (2003) Unifying theory and methodology in biogeography. In:
904 *Evolutionary Biology*. Springer, pp. 1-25.
- 905 Lieberman BS, Eldredge N (1996) Trilobite biogeography in the Middle Devonian:
906 geological processes and analytical methods. *Paleobiology*, 66-79.
- 907 Lieberman BS, Miller III W, Eldredge N (2007) Paleontological patterns,
908 macroecological dynamics and the evolutionary process. *Evolutionary Biology*, **34**, 28-48.

- 909 Liow LH, Reitan T, Harnik PG (2015) Ecological interactions on macroevolutionary
 910 time scales: clams and brachiopods are more than ships that pass in the night. *Ecology*
 911 *letters*, **18**, 1030-1039.
- 912 Liow LH, Uyeda J, Hunt G (2022) Cross-disciplinary information for understanding
 913 macroevolution. *Trends in ecology & evolution*.
- 914 Lovejoy, S. 2013. What is climate? EOS 94, (1), 1 January:p1-2.
- 915 Lovejoy, S. 2015. A voyage through scales, a missing quadrillion and why the climate
 916 is not what you expect. *Climate Dyn.* 44:3187-3210.
- 917 Lovejoy, S. 2019. Weather, Macroweather and Climate: our random yet predictable
 918 atmosphere. Oxford U. Press, New York, N.Y. USA
- 919 Lovejoy, S. 2022. Fractional relaxation noises, motions and the fractional energy
 920 balance equation. *Nonlinear Proc. in Geophys.* in press.
- 921 Lovejoy, S., W. J. C. Currie, Y. Tessier, M. Claeredeboudt, J. Roff, E. Bourget, and D.
 922 Schertzer. 2001. Universal Multifractals and Ocean patchiness Phytoplankton, physical
 923 fields and coastal heterogeneity. *J. Plankton Res.* 23:117-141.
- 924 Mandelbrot, B. B., and J. W. Van Ness. 1968. Fractional Brownian motions, fractional
 925 noises and applications. *SIAM Review* 10:422-450.
- 926 Mayhew PJ, Bell MA, Benton TG, MCGowan AJ (2012) Biodiversity tracks
 927 temperature over time. *Proceedings of the National Academy of Sciences*, **109**, 15141-
 928 15145.
- 929 Markov AV, Korotayev AV (2007) Phanerozoic marine biodiversity follows a
 930 hyperbolic trend. *Palaeoworld*, **16**, 311-318.
- 931 Marshall LG, Webb SD, Sepkoski JJ, Raup DM (1982) Mammalian evolution and the
 932 great American interchange. *Science*, **215**, 1351-1357.

933 Mathes GH, Van Dijk J, Kiessling W, Steinbauer MJ (2021) Extinction risk controlled
 934 by interaction of long-term and short-term climate change. *Nature Ecology & Evolution*,
 935 **5**, 304-310.

936 Mcinerney FA, Wing SL (2011) The Paleocene-Eocene Thermal Maximum: A
 937 perturbation of carbon cycle, climate, and biosphere with implications for the future.
 938 *Annual Review of Earth and Planetary Sciences*.

939 Meyers SR, Siewert SE, Singer BS, Sageman BB, Condon DJ, Obradovich JD, Jicha BR,
 940 Sawyer DA (2012) Intercalibration of radioisotopic and astrochronologic time scales for
 941 the Cenomanian-Turonian boundary interval, Western Interior Basin, USA. *Geology*, **40**,
 942 7-10.

943 Melott A, Bambach R (2014) Analysis of periodicity of extinction using the 2012
 944 geological timescale. *Paleobiology*, **40**, 176-195.

945 Miller A, Foote M (2009) Epicontinental seas versus open-ocean settings: the
 946 kinetics of mass extinction and origination. *Science*, **326**, 1106-1109.

947 Miller, K. S., and B. Ross. 1993. An introduction to the fractional calculus and
 948 fractional differential equations. John Wiley and Sons, New York.

949 Mitchell EG, Harris S, Kenchington CG, Vixseboxse P, Roberts L, Clark C, Dennis A,
 950 Liu AG, Wilby PR (2019) The importance of neutral over niche processes in structuring
 951 Ediacaran early animal communities. *Ecology letters*, **22**, 2028-2038.

952 Nee S (2006) Birth-death models in macroevolution. *Annual Review of Ecology*,
 953 *Evolution, and Systematics*, 1-17.

954 Newman M, Palmer R (2003) *Modeling extinction*, Oxford University Press, Oxford.

955 Newman ME (1997) A model of mass extinction. *Journal of Theoretical Biology*, **189**,
 956 235-252.

- 957 O'brien CL, Robinson SA, Pancost RD, Damsté JSS, Schouten S, Lunt DJ, Alsenz H,
 958 Bornemann A, Bottini C, Brassell SC (2017) Cretaceous sea-surface temperature
 959 evolution: Constraints from TEX86 and planktonic foraminiferal oxygen isotopes. *Earth-*
 960 *Science Reviews*, **172**, 224-247.
- 961 Payne JL, Finnegan S (2007) The effect of geographic range on extinction risk during
 962 background and mass extinction. *Proceedings of the National Academy of Sciences*, **104**,
 963 10506-10511.
- 964 Podlubny, I. 1999. Fractional Differential Equations. Academic Press, San Diego,
 965 United States.
- 966 Procyk, R., S. Lovejoy, and R. Hébert. 2022. The Fractional Energy Balance Equation
 967 for Climate projections through 2100. *Earth Syst. Dynam.* 13:81–107.
- 968 Raup D (1992a) *Extinction: bad genes or bad luck?*, WW Norton & Company, New
 969 York.
- 970 Raup DM (1985) Mathematical models of cladogenesis. *Paleobiology*, **11**, 42-52.
- 971 Raup DM (1991) A kill curve for Phanerozoic marine species. *Paleobiology*, **17**, 37-
 972 48.
- 973 Raup DM (1992b) Large-body impact and extinction in the Phanerozoic.
 974 *Paleobiology*, **18**, 80-88.
- 975 Raup DM (1994) The role of extinction in evolution. *Proceedings of the National*
 976 *Academy of Sciences*, **91**, 6758-6763.
- 977 Raup DM, Valentine JW (1983) Multiple origins of life. *Proceedings of the National*
 978 *Academy of Sciences*, **80**, 2981-2984.
- 979 Reitan T, Liow LH (2017) An unknown phanerozoic driver of brachiopod extinction
 980 rates unveiled by multivariate linear stochastic differential equations. *Paleobiology*, **43**,
 981 537-549.

- 982 Roopnarine PD (2003) Analysis of rates of morphologic evolution. *Annual Review of*
983 *Ecology, Evolution, and Systematics*, 605-632.
- 984 Sadler PM (1981) Sediment accumulation rates and the completeness of
985 stratigraphic sections. *The Journal of Geology*, 569-584.
- 986 Saupe E, Qiao H, Donnadieu Y, Farnsworth A, Kennedy-Asser A, Ladant J, Lunt D,
987 Pohl A, Valdes P, Finnegan P (2019) Extinction intensity during Ordovician and Cenozoic
988 glaciations explained by cooling and palaeogeography. *Nature Geoscience*.
- 989 Schopf TJM (1979) Evolving paleontological views on deterministic and stochastic
990 approaches. *Paleobiology*, **5**, 337-352.
- 991 Sepkoski JJ (1984) A kinetic model of Phanerozoic taxonomic diversity. III. Post-
992 Paleozoic families and mass extinctions. *Paleobiology*, **10**, 246-267.
- 993 Sepkoski JJJ (1996) Competition in macroevolution: the double wedge revisited. In:
994 *Evolutionary paleobiology*. University of Chicago Press, Chicago (eds Jablonski D, Erwin
995 DH, Lipps JH). The University of Chicago Press, Chicago, pp. 211-255.
- 996 Song H, Wignall PB, Song H, Dai X, Chu D (2019) Seawater temperature and
997 dissolved oxygen over the past 500 million years. *Journal of Earth Science*, **30**, 236-243.
- 998 Spiridonov A, Balakauskas L, Lovejoy S (2022) Longitudinal expansion fitness of
999 brachiopod genera controlled by the Wilson cycle. *Global and Planetary Change*, 103926.
- 1000 Spiridonov A, Brazauskas A, Radzevičius S (2015) The role of temporal abundance
1001 structure and habitat preferences in the survival of conodonts during the mid-early
1002 Silurian Ireviken mass extinction event. *PLoS ONE*, **10**, e0124146.
- 1003 Spiridonov A, Brazauskas A, Radzevičius S (2016) Dynamics of abundance of the
1004 mid- to late Pridoli conodonts from the eastern part of the Silurian Baltic Basin:
1005 multifractals, state shifts, and oscillations. *American Journal of Science*, **316**, 363-400.

- 1006 Spiridonov A, Kaminskas D, Brazauskas A, Radzevičius S (2017a) Time hierarchical
1007 analysis of the conodont paleocommunities and environmental change before and during
1008 the onset of the lower Silurian Mulde bioevent – A preliminary report. *Global and*
1009 *Planetary Change*, **157**, 153-164.
- 1010 Spiridonov A, Lovejoy S (2022) Life rather than climate influences diversity at scales
1011 greater than 40 million years. *Nature*, **607**, 307–312.
- 1012 Spiridonov A, Samsonė J, Brazauskas A, Stankevič R, Meidla T, Ainsaar L,
1013 Radzevičius S (2020a) Quantifying the community turnover of the uppermost Wenlock
1014 and Ludlow (Silurian) conodonts in the Baltic Basin. *Palaeogeography, Palaeoclimatology,*
1015 *Palaeoecology* **549**, 109128.
- 1016 Spiridonov A, Stankevič R, Gečas T, Brazauskas A, Kaminskas D, Musteikis P,
1017 Kaveckas T, Meidla T, Bičkauskas G, Ainsaar L, Radzevičius S (2020b) Ultra-high
1018 resolution multivariate record and multiscale causal analysis of Pridoli (late Silurian):
1019 implications for global stratigraphy, turnover events, and climate-biota interactions.
1020 *Gondwana Research*, **86**, 222-249.
- 1021 Spiridonov A, Stankevič R, Gečas T, Šilinskas T, Brazauskas A, Meidla T, Ainsaar L,
1022 Musteikis P, Radzevičius S (2017b) Integrated record of Ludlow (Upper Silurian) oceanic
1023 geobioevents–Coordination of changes in conodont, and brachiopod faunas, and stable
1024 isotopes. *Gondwana Research*, **51**, 272-288.
- 1025 Stanley SM (1979) *Macroevolution: Pattern and Process*, W. H. Freeman, San
1026 Francisco.
- 1027 Tomašových A, Jablonski D, Berke SK, Krug AZ, Valentine JW (2015) Nonlinear
1028 thermal gradients shape broad-scale patterns in geographic range size and can reverse
1029 Rapoport's rule. *Global Ecology and Biogeography*, **24**, 157-167.

- 1030 Vakulenko, S. A., I. Sudakov, and L. Mander. 2018 The influence of environmental
 1031 forcing on biodiversity and extinction in a resource competition model. *Chaos*
 1032 28, :031101
- 1033 Van Dam JA, Aziz HA, Sierra MÁÁ, Hilgen FJ, Van Den Hoek Ostende LW, Lourens LJ,
 1034 Mein P, Van Der Meulen AJ, Pelaez-Campomanes P (2006) Long-period astronomical
 1035 forcing of mammal turnover. *Nature*, **443**, 687-691.
- 1036 Van Valen L (1973) A new evolutionary law. *Evolutionary Theory*, **1**, 1-30.
- 1037 Veizer J, Ala D, Azmy K, Bruckschen P, Buhl D, Bruhn F, Carden GA, Diener A, Ebneh
 1038 S, Godderis Y (1999) $^{87}\text{Sr}/^{86}\text{Sr}$, $\delta^{13}\text{C}$ and $\delta^{18}\text{O}$ evolution of Phanerozoic seawater. *Chemical*
 1039 *Geology*, **161**, 59-88.
- 1040 Veizer, J., Y. Godderis, and L. M. Francois. 2000. Evidence for decoupling of
 1041 atmospheric CO₂ and global climate during the Phanerozoic eon. *Nature*, 408:698–701.
- 1042 Venckutė-Aleksienė A, Spiridonov A, Garbaras A, Radzevičius S (2018) Integrated
 1043 foraminifera and $\delta^{13}\text{C}$ stratigraphy across the Cenomanian–Turonian event interval in
 1044 the eastern Baltic (Lithuania). *Swiss Journal of Geosciences*, **111**, 341-352.
- 1045 Vermeij GJ (1977) The Mesozoic marine revolution: evidence from snails, predators
 1046 and grazers. *Paleobiology*, **3**, 245-258.
- 1047 Vermeij GJ (2019) Power, competition, and the nature of history. *Paleobiology*, **45**,
 1048 517-530.
- 1049 Vrba ES (1985) Environment and evolution: alternative causes of the temporal
 1050 distribution of evolutionary events. *South African Journal of Science*, **81**, 229-236.
- 1051 Vrba ES (1993) Turnover-pulses, the Red Queen, and related topics. *American*
 1052 *Journal of Science*, **293**, 418-452.
- 1053 Zachos, J., M. Pagani, L. Sloan, E. Thomas, and K. Billups. 2001. Trends, Rhythms, and
 1054 Aberrations in Global Climate 65 Ma to Present. . *Science* 292(5517):686-693.

- 1055 Žliobaitė I (2022) Recommender systems for fossil community distribution
1056 modelling. *Methods in Ecology and Evolution*, **13**, 1690-1706.
- 1057 Žliobaitė I, Fortelius M, Stenseth NC (2017) Reconciling taxon senescence with the
1058 Red Queen's hypothesis. *Nature*, **552**, 92-95.
- 1059

1060 **Appendix A: Numerical simulations**

1061 Since the model is linear, the obvious simulation method is to use Fourier techniques.
1062 The main problem is that the small scales have singular power law filters that are not trivial to
1063 discretize, there may also be some long time (low frequency) issues. A convenient way is to
1064 use techniques developed for simulating ffRn processes discussed in (Lovejoy 2022). ffRn
1065 processes can be simulated by convolving Gaussian white noises with the ffRn Green's
1066 function $G_{\alpha,h}$ (eqs. 9,10). A somewhat better numerical technique is to use the step response
1067 Green's function ($=G_{\alpha+1,h}$ it is the smoother – and hence easier to handle integral of $G_{\alpha,h}$),
1068 followed by a numerical differentiation.

1069

1070

1071 **Figure Captions**

1072 Fig. 1: The impulse (delta function) response $G_{\alpha,0}(t) = t^{\alpha-1} / \Gamma(\alpha)$ for fractional integrals of
 1073 order α normalized for the same response after 1 Myr. The bottom corresponds to the turnover
 1074 (E_+) response $\alpha = 1/4$ and the top corresponds to the temperature (T) response with $\alpha = 3/4$.
 1075 Notice the long term effects.

1076 Fig. 2: The impulse response $G_{\alpha,h}(t/\tau)$, with $\alpha = 1/4$, $h = 1/2$ corresponding to the
 1077 diversity(D) response, for critical transition times $\tau = 1, 4, 16, 64, 256$ Myrs (bottom to top).
 1078 The empirical value is $\tau \approx 40$ Myrs (SL).

1079 Fig. 3: This shows the Phanerozoic marine animal macroevolutionary analysis of the 6
 1080 series discussed in this paper; D, T, O, E are replotted from SL. The dashed lines show the
 1081 theory slopes (eq. 44) with transition at $\Delta t \approx 40$ My i.e. $\log_{10}\Delta t \approx 1.6$.

1082 Fig. 4: The (normalized) pairwise correlations of the 15 pairs of the 6 series as functions
 1083 of lag. Several of these are reproduced from SL.

1084 Fig. 5: The previous 2^{14} simulation degraded from $1/4$ Myr resolution to 1 Myr. Curves
 1085 normalized by their standard deviations and then offset by 5 units in the vertical for clarity.

1086 Simulation $2^{14} = 16384$ points with theoretical slopes indicated. The transition scale τ is
 1087 $2^7 = 128$ units, indicated by dashed vertical lines. If the model was at 250kyr resolution, the
 1088 cross over is at 32Myrs, the length of the simulation is: 4 Gyrs. Parameters $\alpha = 0.25$, $h = 0.5$,
 1089 $\rho_E = \rho_{TE} = 0.5$, $\rho_D = \rho_{TD} = -0.1$ (with derived DE correlation $\rho_{DE} = -0.9$).

1090 Fig 6: Simulation $2^{14} = 16384$ points with theoretical slopes indicated. The transition
 1091 scale τ is $2^7 = 128$ units, indicated by dashed vertical lines. If the model was at 250kyr resolution,
 1092 the cross over is at 32 Myrs, the length of the simulation is: 4 Gyrs. Parameters $\alpha = 0.25$, $h =$
 1093 0.5 , $\rho_E = \rho_{TE} = 0.5$, $\rho_D = \rho_{TD} = -0.1$ (with derived DE correlation $\rho_{DE} = -0.9$).

1094 Fig. 7: The 15 pairwise correlations from the 2^{14} realization above. Only two of the
 1095 correlations were prescribed and this, only at a single resolution, the rest are consequences of
 1096 the model, the two exponents a , h and the cross-over time $\tau = 2^7$ (shown as short dashed vertical
 1097 lines). The two prescribed correlations (DT, TE₊) are shown as solid horizontal lines, and the
 1098 derived correlations (DE₊ from DT, TE₊, eq. 32) and then TE, TO (predicted to be equal to
 1099 equal to TE₊ at long lags – eq. 39) and DE, DO (predicted to be equal to DE₊, at long lags see
 1100 eq. 39). Note that these are from a single realization of the process not the ensemble average.
 1101 In addition, the statistics of some are fairly sensitive to irregularly sampled (and small size) of
 1102 the empirical data, compare with fig. 10 below.

1103 Fig. 8: Model - simulation comparison with all series normalized by their standard
 1104 deviations. The simulation was at 1Myr resolution and the sampled at the same (irregular)
 1105 times as the data (84 points over the last 500Myrs). Each curve was displaced by 5 units in the
 1106 vertical for clarity. Due to causality, the series are asymmetric with time running from right to
 1107 left. The simulation is on the right.

1108 Fig. 9: From the 1Myr resolution simulations discussed above (Brown) and in fig. 8 and
 1109 resampled at the data times (red), black is data. The relative vertical offsets of the curves are
 1110 not significant, they correspond to specific normalizations / nondimensionalizations. We see
 1111 that in general, the resampling at the data times (red) yields a closer fit to the data (black) than
 1112 the analysis of the simulation at uniform (1Myr) intervals, this is especially true for E_- , O , E ,
 1113 E_+ .

1114 Fig. 10: Same simulation as above, compared with data (black). Brown is a uniform
 1115 1Myr resolution, red is the simulation resampled at the data times. The resampling notably
 1116 improves the correlations for DE_+ , DO , DE , E_+E_- , $E-O$, OE , and to a lesser extent the OE , TE_+
 1117 comparisons for the others it is about the same.